Brazo's River Authority

May 2003 CHIRP Sub-bottom Profiler Survey and Sediment Core Analyses of:

Aquilla Lake Granger Lake Limestone Lake Lake Proctor

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BRA-CHIRP Sub-bottom Profiler survey

Scope of work:

The goal of this project was to collect seismic data to determine the thickness of sediment within four Brazos River Authority reservoirs. To accomplish this, a CHIRP sub-bottom profiler (see description below) was used to collect high-resolution seismic profiles of the lakebeds. This equipment is designed to be used for subsurface geologic investigations of shallow (0-100 m) subsurface sediments in aquatic settings. Navigation was provided via differential GPS.

This work was carried out in conjunction with a team from Baylor University who collected core samples for calibrating and verifying a multi-frequency depth sounder. Additional core material gathered was x-rayed and profiled for ²¹⁰Pb and ¹³⁷Cs to provide an estimation of sedimentation rates and sedimentation histories in these systems.

Materials and Methods

Core Analysis:

Water Content

Samples were placed into aluminum dishes after extrusion, weighed to the ten thousandth decimal place of a gram (g) and placed into ovens for at least 24 hrs. The samples were removed and weighed again to obtain water content. The porosities were calculated to obtain corrected depths for ²¹⁰Pb calculations.

Grain size

Grain size distributions were determined for the cores following the Folk (1980) methodology. First, sediment samples weighing approximately 15 g were separated with Calgon (Sodium Metahexaphoshate Soap) causing deflocculation, followed with wet sieving the samples with deionized water into a 1000 ml graduated cylinder. Next the graduated cylinders were filled with deionized water and the sand content from the sieves were placed into an aluminum dish and weighed. Graduated cylinder samples were plunged for 20 seconds and left undisturbed for an additional 20 seconds, and then immediately a 20 ml pipette withdraws (4 phi) were taken at a depth of 10 cm. Approximately 2 hrs later, depending on air temperature, another 20 ml pipette withdraws (8 phi) were taken at a depth of 20 cm. All samples were dried, weighed and placed into a spreadsheet to determine fine to coarse grain distribution.

¹³⁷Cs and ²¹⁰Pb Analysis

Samples for geochemical analyses were removed from homogeneous sample bags to determine the distribution of ¹³⁷Cs and ²¹⁰Pb. Sediment samples for ¹³⁷Cs analysis were wet packed and sealed with electrical tape in 60X15 mm Petri dish. The samples were counted on Canberra 2000 mm² planar coaxial detectors for 1-2 days per sample.

Gamma energy activities were measured for ¹³⁷Cs at the decay energy signal of 661.7 keV.

The excess activity of ²¹⁰Pb was measured by alpha spectroscopy following the methodology by Santschi et al. (1980, 1999, 2001). Sediment samples weighing approximately 15 g were dried for each selected depth, then pulverized and homogenized with a mortar and pestle. Approximately 0.5 g aliquots were placed into 100 ml Teflon beakers and leached with 15 ml of HCL and HNO₃, and 10 ml of HF. The samples were spiked with 500 µl of ²⁰⁹Po and baked to near dryness on hotplates. Then 15 ml of HCL and HNO₃ were added to the samples and taken to near dryness again. The Teflon beakers were rinsed with 10 ml of HCL and baked to complete dryness. Samples were diluted with 50 ml of 1.5 N of HCL, and ascorbic acid was stirred in with a magnetic stir bars until the samples turned to clear color appearance. A 1-cm² silver planchet was placed opposite of the magnetic stir bar. The beakers were covered with watch glasses and heated approximately to 80 degrees for 2.5 hrs. Silver plates were removed and counted for 1-2 days for alpha decay on Canberra Quad Alpha Spectrometer connected to S100 multi-channel analyzer to reach a counting error of ~ 2 % or less. Supported activities were estimated from total ²¹⁰Pb values deep in cores where excess activities have decayed to negligible values, and were subtracted from the total activity to determine excess activities (Dellapenna et al. 2003).

X-Radiography

The cores were split and half of the core was transferred to a plexiglass tray for x-raying. There was considerable effort in transferring the cores from the core barrels because the sediment tended to adhere to the core barrel and as a result, there was some disturbance to the sediment. X-radiographs were taken with a portable Medison X-radiograph Unit model PX-15HF at 72 kV and 150 mAs for each plexiglass tray. Fuji sheet negatives were developed and scanned with Microtek ScanMaker 9600XL into digital forms.

Equipment:

CHIRP Sub-bottom profiler - The CHIRP sub-bottom profiler sends out a "chirp" of energy from 2-16 kHz. The return of this signal to a receiver is used to create a linear cross section of the bay bottom along the track of the tow vessel. Return signals differ based on the bottom type encountered, some sound bouncing off the bottom itself, whereas some penetrates and reflects off interfaces within the sediment. Repeating the send/receive action, a strip chart makes a profile showing the penetration and the characteristics of the reflections. The sequence of subsurface sediments can be resolved on the decimeter scale and the sub-bottom profiler can typically penetrate between 10-100 m into the seabed.

Navigation - Position was determined using a Trimble AG132 Trimble Differential GPS with Everest, which was fed digitally into a Gateway Laptop and logged using Hypack Coastal Oceanographic software. Hypack was also used to create survey lines and was used as a chart plotter for navigation of these lines. The layback of the tow fish (fore and aft) to the DGPS antennae was estimated by measuring the distance from the antennae to

the transom and then measuring the amount and angle of the tow cable from the transom to the tow fish. The offset (a beam) was determined by measuring the offset of the DGPS antennae to the tow point of the cable on the davit. Errors for position of the fish in relation to the bottom should be limited to the error inherent in DGPS (less than 1 meter).

Results

Field work to collect the CHIRP lines were conducted off of the TAMUG Coastal Geology research boat, the R/V Cavalla (Figure 1), which is a Parker 25 using an Edgetech 272 CHIRP (Figure 2). Granger Lake was surveyed on 05/07/03, Limestone Lake was surveyed on 05/08/03, Aguilla and Proctor Lakes were surveyed on 5/09/03. Appendix A contains all of the CHIRP Lines collected for each Lake. The thickness of sediment at each point where it could be determined is shown on each line. In most lines, the pre-impoundment surface is a hard, dense surface, with reservoir sediment being soft, high watercontent material above the pre-impoundment surface. As a result, the preimpoundment surface generally appears to be a dark reflector and the surface of the new sediment appears as a faint line which defines the sediment-water interface (mudline). The reservoir sediment is composed of mud with a high water content and appears in the CHIRP record as a faint layer. Because these reservoirs are sediment filled incised valleys, the pre-impoundment surface is often highly irregular, with reservoir sediment generally filling pre-impoundment topographic lows. In each reservoir, there is only one reflector at depth, which represents the pre-impoundment surface, with the faint layer of reservoir sediment above. For each line, the thickness of reservoir sediment is noted at key points along the line. Because of the irregular nature of the pre-impoundment surface, no attempts were made to create an isopach map of reservoir sediment.

Granger Lake - Lines G1-G5 are the survey lines for Granger Lake, it was not decided until after Granger Lake was surveyed to add additional tie lines to the surveys.

Limestone Lake - The survey in Limestone Lake was conducted between the dam and a bridge, the R/V Cavalla would not fit under the bridge at the time of the survey, and no boat ramps were found above the bridge. Limestone Lake Lines 1-6 are cross lake lines, Line 7 runs down the axis of the lake.

Aquilla Lake - Lines A1 runs across the back of the dam. Lines A2-A4 are diagonal across the main portion of the lake.

Proctor Lake - Lines P1-P5 are cross lake lines. Lines P6 and P7 were not run because there were too many trees sticking out of the water. Line P8 is a tie line, it was not run between P3 and P5 because of navigational obstructions caused by trees in the lake.

Sediment Core Data

All cores were collected by the Baylor University group and core analyses were made at the request of this group. The cores were collected in aluminum tubing, which proved to be problematic because we could not x-ray through the tubes. The high clay content of the sediment resulted in strong adhesion of the sediment to the core barrels when the cores were split.

X-radiographs - The x-radiographs from the cores for this project were of poor quality. Part of the problem is the small diameter of the core barrels used. The sediment has a high clay content, and with narrow core barrels, surface friction between the core barrel and the sediment is much more difficult to overcome, as a result, it appears that the sediment balled up in the core barrel. In addition, when the cores were split, the adhesion problem made transferring the sediment from the core barrel to the x-ray tray difficult as well. Consequently, the x-rays were not nearly as useful as originally hoped.

Water Content Data - Water content data is used in the ²¹⁰Pb calculations. Water content is a function of both the grain size of the sediment and is much higher in recently deposited sediments than older, more compacted sediments. As a result, it is useful as a stratigraphic indicator and often mimics the radioisotope profiles. Water content profiles are provided for each core.

Grain size Data - Sand, silt, clay distribution was determined for 1 cm thick intervals every 5 cm throughout the cores. In general, the sediment in the reservoirs were clay dominated with little sand. The sand content is equivalent to one minus the mud content.

Radioisotope Data - ²¹⁰Pb radioisotope profiles show that for each core, there appears to be steady-state accumulation, with little or no seabed mixing. Because these cores were of reservoir sediment and the reservoirs are less than 50 years old, the ²¹⁰Pb in the sediment are all at excess levels. Accumulation rates were determined by using the regression line of the excess ²¹⁰Pb activities of the cores. In each case the calculation was quite straight forward. ¹³⁷Cs was found at the base of each core, in the last sample of reservoir sediment. ¹³⁷Cs is a bomb produced isotope first measurable in the environment in 1954 with atmospheric testing of nuclear bombs. These reservoirs were created after the first appearance of ¹³⁷Cs and all of the reservoir sediment which was introduced after 1954 contains ¹³⁷Cs.

Aquilla Lake Core 2- (Figures 3-6) - This core generally had over 90% clay, as a result, transferring the core to the x-ray tray resulted in significant core damage resulting in an extremely poor x-radiograph. The water content profile shows little change from the top to the bottom of the core. The ²¹⁰Pb profile shows a nearly vertical profile as well. A best fit line for what may be a decay portion of the core was used to estimate accumulation rates, which are estimated to be 0.49 cm/yr. These rates should be used with caution, because there really is not much of a decay profile for this core. ¹³⁷Cs was present at the base of the core, suggesting that the sediment at the base of the core maybe

as old as 40 y. An accumulation rate of 2.5 cm/y would be consistent with this observation. The grain size profiles show two layers which are silt rich, one at 15 cm and another at 85 cm, these may represent flood layers or reservoir discharge periods. Where silt content fluctuates, interpretations of ²¹⁰Pb profiles becomes problematic. This core was taken in a narrow, deep channel.

Aquilla Lake Core 2B (Figures 7-10)- The x-radiograph shows the core to have a light layer at 20 cm, which correlates to the grain size profile, which shows this layer to contain a 5% increase in sand content. We had a box of x-ray film where each piece of film appears to have been exposed on the edges, which is why there are light bands on each side of the film. Unfortunately, this was not discovered until after the cores had been processed. The x-radiograph also contains a number of light spots which probably are the result of methane bubbles. These are quite common in marine sediment cores. There also appears to be a hiatal surface at 45 cm, which correlates to an increase in silt content. The ²¹⁰Pb profile suggests that there is a steady state decrease in isotope activity down to 30 cm, then the activity becomes linear. The water content profile shows a similar trend, with an abrupt decrease in water content at 30 cm and a large fluctuation in silt content between 30-50 cm. Because of the dramatically higher surface area to volume of clay versus silt, an increase in silt content will decrease ²¹⁰Pb activities. Presumably this core was taken proximal to Core 2, we have no position for Core 2B.

Aquilla Lake Core 3B (Figures 11-14)- The x-radiograph from this core is also of low quality, not sure what happened with it. There is a 10 cm thick mud layer at the surface of this core and the sand content progressively increases with depth, except for a high silt content layer at 40 cm. The ²¹⁰Pb activity decrease with depth may be more a function of an increase in sand content than a decay of ²¹⁰Pb. The rates estimated for this core should be used with caution. This core was collected near shore, where transport of sand from the shore line would be expected. Also, fluctuations in lake level would move the shore closer to the core site when lake level was lower. The periods of high sand content my represent lower lake levels. Alternatively, these sand layers may represent flood deposits.

Limestone Lake Cores

We were not provided with core locations for any of the cores we have for Limestone Lake, so interpretations were not made based on the location of the cores.

Limestone Lake LS-2B (Figures 15-18)- This core is only 20 cm long. The grain size profile shows that the sediment becomes muddier with depth. The shape of the ²¹⁰Pb profile probably reflects this textural change more than anything. The x-radiograph shows there to be a distinct layer at the surface and mottled mud with clay balls within the rest of the core.

Limestone Lake LS-3 (Figures 19-22)- There is a great deal of cyclicity to the silt and clay content in this core, suggesting a strong flood signal. The sand content in the lower 30 cm of this core ranges from 40% at the base to $\sim 10\%$ at the top of the interval. This

fining upwards sequence is probably the result of sand being transported as bedload as the reservoir is filling and once it has filled, little bedload transport is occurring. There is a degree of variability in the ²¹⁰Pb profiles which probably results from variations in sediment textures, however, the overall trend suggests steady-state accumulation. The position of the core location, in the middle of the reservoir suggests that the alternating clay rich and clay poor layers is probably results from high discharge/flood events, during high sediment input, first the silt will settle out then the clays will settle. This site could probably be viewed as a site where the distal signal of a flood is recorded. That is to say, it is further removed from the sources of coarse sediment so all that is seen is the settling of fines

Limestone Lake LS-4 (Figures 23-25)- This core is only 10 cm long. It is 95% mud. There was not much reservoir fill at this site.

Lake Proctor 1B (Figures 26-29)- The high water content in the upper 5 cm, along with the uniform ²¹⁰Pb activity and the uniform grain size for this interval suggests this is a flood layer. The location of this core so close to shore and also proximal to the head of the lake would suggest that coarser sediment would be delivered to this site during high discharge events.

Granger Lake 1B (Figures 30-33)- The extreme range in water content as well as the fining upwards sequence shown in the grain size profiles suggest that the base of this core is in the pre-impoundment surface at about 16 cm. This core was taken directly behind the reservoir and it would appear that the changes in the sediment textures reflects the history of sediment fill and changes in the flow regime as the water depth increases. At the base of the core, the sediment is much coarser, water depths are shallower, the cross sectional area of the river would be greater, and the critical shear stress on the river bed would be high enough to allow for bedload transport of coarse material. As the water depth increases, the cross sectional area of the river increases, the flow strength decreases and the less coarse material will be transported as bedload transport becomes less. As a result, the sediment textures will become much finer.

Discussion and Conclusion

The CHIRP data shows that the each reservoir is a filled incised valley. The incised valleys were generally flat just up stream of the reservoirs, and the pre-impoundment surface is quite irregular upstream from the dam. Because the pre-impoundment surface is so irregular, no effort was made to make an isopach map of reservoir fill. Where feasible, tie lines were run down the axis of the reservoir so that correlations could be made between cross lake CHIRP lines. In each CHIRP line, there was one hard reflector, which represents the pre-impoundment surface. Above this reflector, there is a faint layer of high water-content mud which represents the reservoir sediment fill. The thickness of this fill is noted on each CHIRP line. The high clay-content of the sediment cores caused there to be a great deal of coring artifacts in the x-rays. Grain size analyses reveal clay content to generally be over 50% in most cores and some cores had intervals with clay

content approaching 100%. Radioisotope profiles reveal steady state accumulation in nearly all of the cores. However, large changes in sediment textures have limited the utility of the radioisotope interpretation in several of the cores. ¹³⁷Cs was present at the base of the reservoir fill in each core. Limited detector time precluded running profiles of ¹³⁷Cs, however, it is doubtful that profiles would have aided in interpretations, especially given the high variability of sediment textures in many of the cores. Some interesting trends in grainsize profiles were observed. For cores directly behind the dam, such as Granger Lake Core 1B, the grain size profile at the base of the core going up core shows a fining upwards sequence which reflects the early history of reservoir fill and a decrease in bedload transport as the cross sectional area of the river increases and the strength of the river decreases. Above this, the cores generally show a fluctuation between high and low clay content, reflecting either high discharge events or water release events from the dam. Because there is generally low turbulence in the deeper portions of the reservoir, fall velocity is the major control on sediment deposition. As a result, after high discharge events, such as floods, which deliver a suspended sediment load, first the silts will deposit then the clays. Alternating layers of silt and clay may reflect this trend. Cores collected in the middle of the reservoir may show a similar trend with the lower core record of transition from coarse to find grained material and then a pronounced fluctuation between silt and clay similar to the cores from the lower reservoir. Upstream cores and cores from locations near-shore, will typically have a record which shows alternating sandy and muddy layers, with sandy layers from either low water periods with shore line erosion or from flood events with a delivery of bedload material.

Appendix A

Lake	Core	Longitude	Latitude
Aquilla	Annual	97.11.258'	31.55.656'
Aquilla	2	97.11.289'	31.55.675'
Aquilla	3	97.12.564'	31.55.196'
Aquilla	4	97.12.968'	31.55.168'
Aquilla	5	97.11.736'	31.55.179'
Aquilla	6	97.12.379'	31.54.348'
Granger	1	97.20.178'	30.42.129'
Granger	2	97.20.312'	30.41.793'
Granger	3	97.21.279'	30.41.829'
Granger	4	97.21.212'	30.41.997'
Granger	5	97.21.484'	30.41.363'
Granger	6	97.21.362'	30.41.101'
Limestone	1	96.18.870'	31.19.718'
Limestone	2	96.21.333'	31.25.235'
Limestone	3	96.20.208'	31.23.781'
Limestone	4	96.20.003'	31.23,863'
Limestone	5	96.19.606'	31.19.746'
Proctor	emissi,	98.30.893'	31.58.697'
Proctor	2	98.29.638'	31.58.427'
Proctor	3	98.29.649'	31.58.306'
Proctor	4	98.28.407'	31.59.216'
Proctor	5	98.28.404'	31.59.228'
Proctor	6	98.28.255'	31.59.197'

NAD83

Appendix B

Location of CHIRP Lines

Aquilla Lake

1)	Beginning:	31°55.731' N	97°11.379' W
	Ending:	31°55.683' N	97°11.219' W
2)	Beginning:	31°55.700' N	97°11.292' W
	Ending:	31°54.011' N	97°12.330' W
3)	Beginning:	31°54.421' N	97°12.737' W
	Ending:	31°54.449' N	97°11.655' W
4)	Beginning:	31°54.136' N	97°11.841' W
	Ending:	31°55.281' N	97°12.854' W

Granger Lake

1)	Beginning:	31°40.389' N	97°22.176' W
	Ending:	30°40.565' N	97°22.551' W
2)	Beginning:	30°41.119' N	97°21.330° W
	Ending:	31°41.434' N	97°21.537' W
3)	Beginning:	30°41.385' N	97°20.459' W
	Ending:	31°41.312' N	97°20.449° W
4)	Beginning:	30°42.694' N	97°20.867' W
	Ending:	30°42.861' N	97°20.766' W
5)	Beginning:	30°42.164' N	97°21.586' W
	Ending:	30°41.798' N	97°21.775' W
6)	Beginning:	30°42.350' N	97°22.883' W
	Ending:	30°41.963' N	97°22.706' W

Appendix B

CHIRP Lines from each Reservoir.

Limestone Lake

1)	Beginning:	31°19.901' N	96°18.738' W
	Ending:	31°19.653' N	97°19.724' W
2)	Beginning:	31°23.228′ N	96°19.210' W
	Ending:	31°22.787′ N	96°19.791' W
3)	Beginning:	31°24.339′ N	96°20.482' W
	Ending:	31°24.510′ N	96°20.369' W
4)	Beginning:	31°24.874′ N	96°21.789' W
	Ending:	31°25.604′ N	97°20.557' W
5)	Beginning:	31°26.951′ N	96°21.097′ W
	Ending:	31°27.000′ N	96°20.983' W
6)	Beginning:	31°26.479′ N	96°22.530' W
	Ending:	31°26.072′ N	96°22.537′ W
7)	Beginning:	31°26.401′ N	96°22.736′ W
	Ending:	31°19.873' N	96°18.854' W

Procter Lake

1) Beginning:	31°58.689' N	98°30.894' W
Ending:	31°58.438' N	98°30.883' W
2) Beginning:	31°58.363' N	98°28.960' W
Ending:	31°58.833' N	98°28.981' W
3) Beginning:	31°59.114' N	98°28.778' W
Ending:	31°59.068' N	98°28.147' W
4) Beginning:	31°59.227' N	98°29.751' W
Ending:	31°59.373′ N	98°29.522' W
5) Beginning:	32°01.010′ N	98°30.419' W
Ending:	32°01.061′ N	98°30.356′ W
6) Beginning: Ending:		
7) Beginning: Ending:		
8) Beginning:	31°59.283' N	98°29.111' W
Ending:	31°58.503' N	98°31.035' W

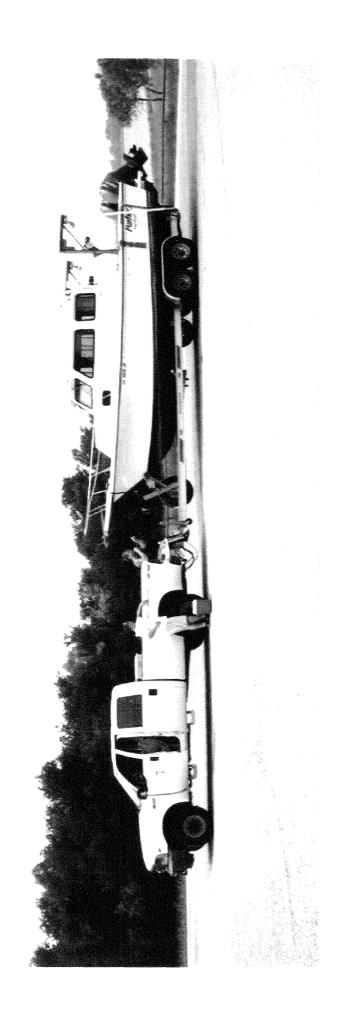
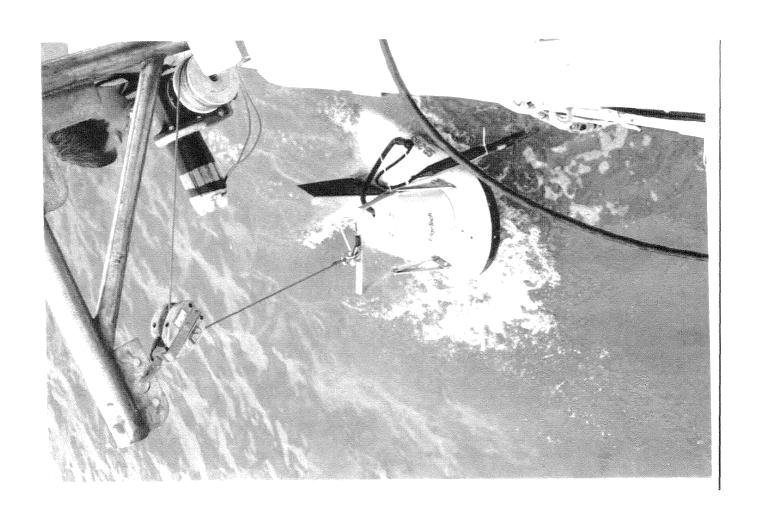


Figure 1- R/V Cavalla rigged and ready for launch in a Brazos River Reservoir

Figure 2- CHIRP being deployed





Aquilla Lake Core 2

Water Content for Aquilla 2 Core

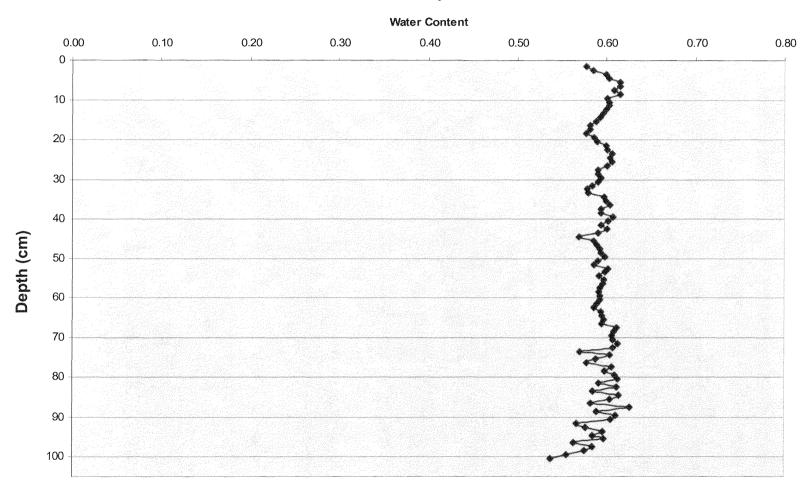


Figure 4

Grain Size Analysis of Brazos River Core Aq 2

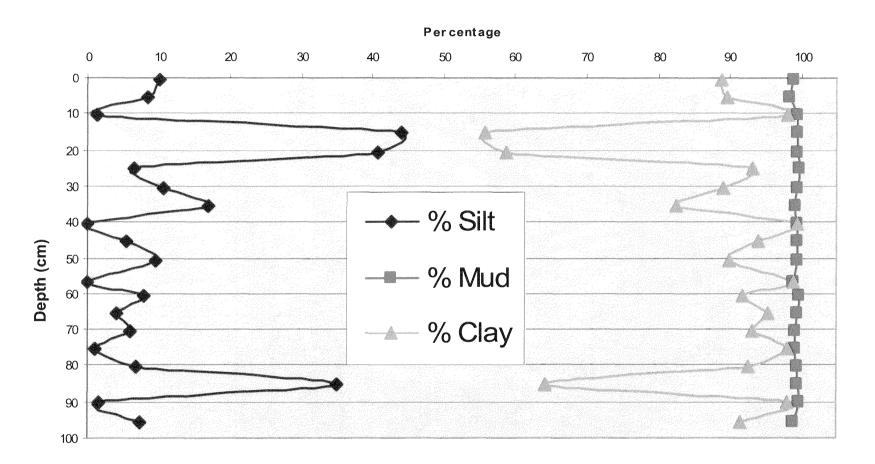


Figure 5

Brazos River - Alpha for Aquilla 2 Core

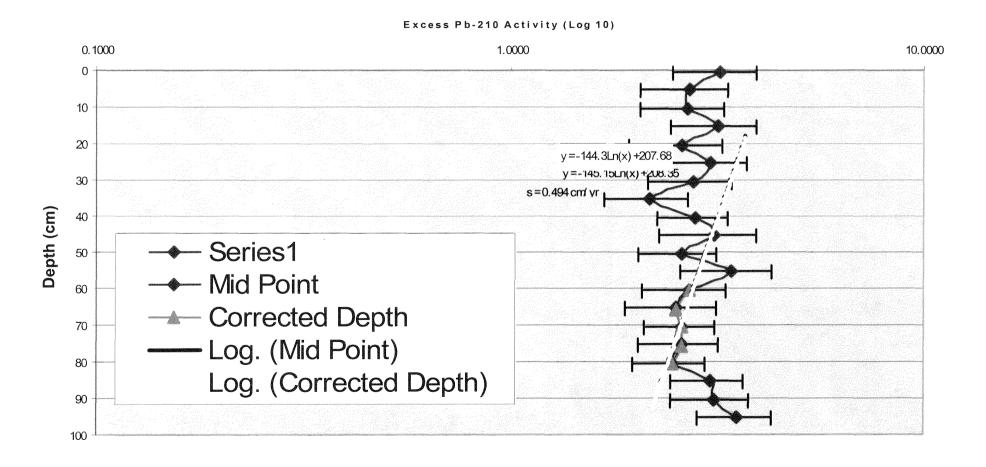
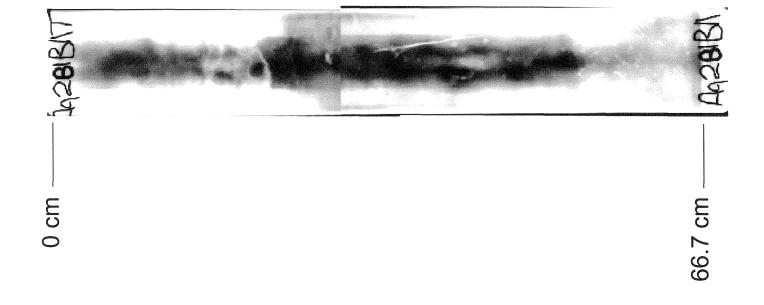


Figure 6



Aquilla Lake Core 2B

Water Content for Aquilla 2-B1 Core

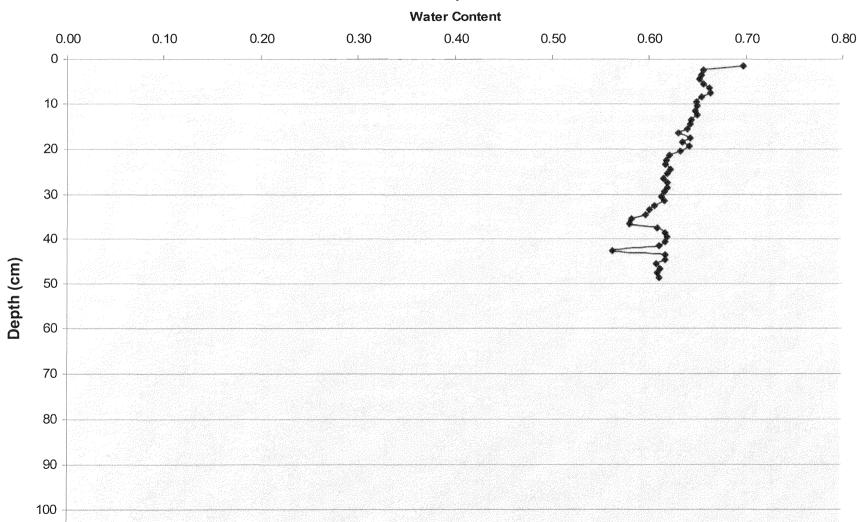


Figure 8

Grain Size Analysis of Brazos River Core Aq 2B1

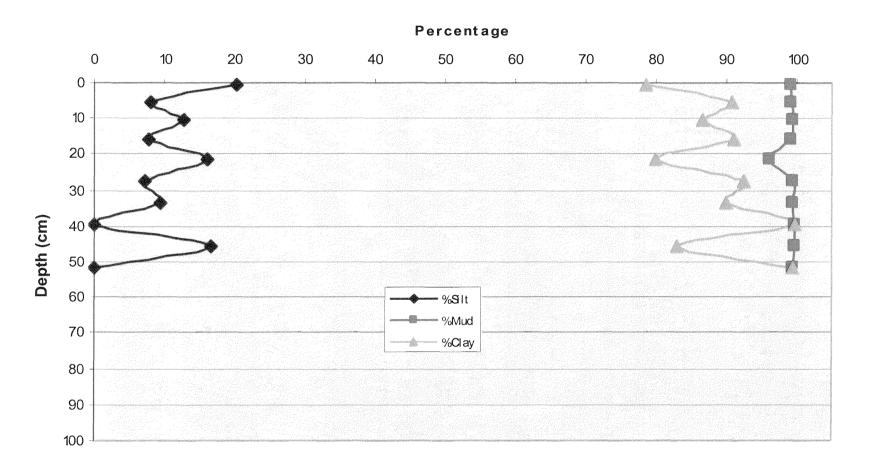


Figure 9

Brazos River - Alpha for Aquilla 2B1 Core

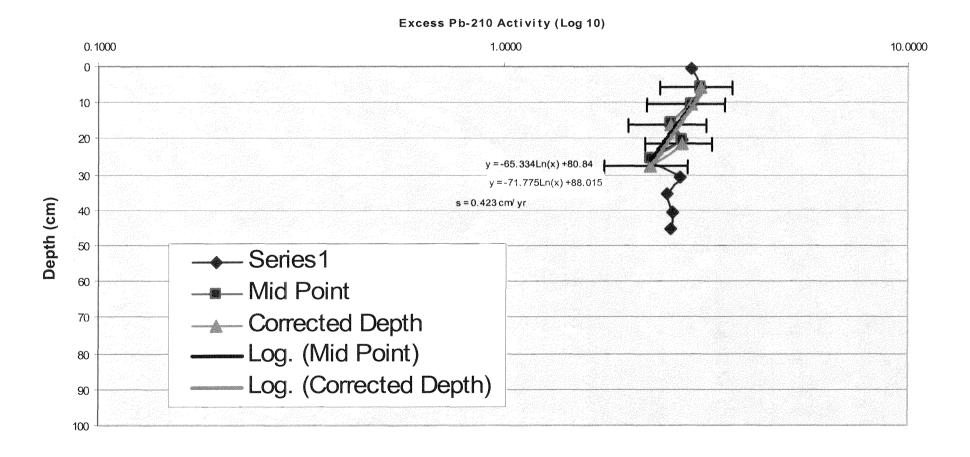
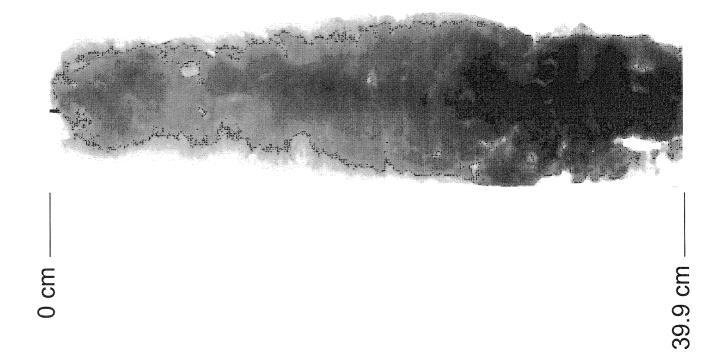


Figure 10



Aquilla Lake Core 3B

Water Content for Aquilla 3B Core

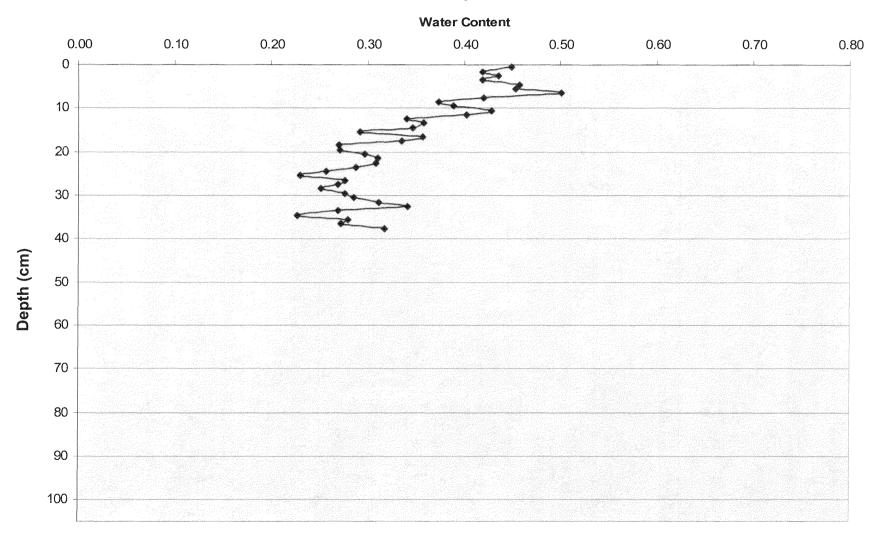


Figure 12

Grain Size Analysis of Brazos River Core Aq 3B

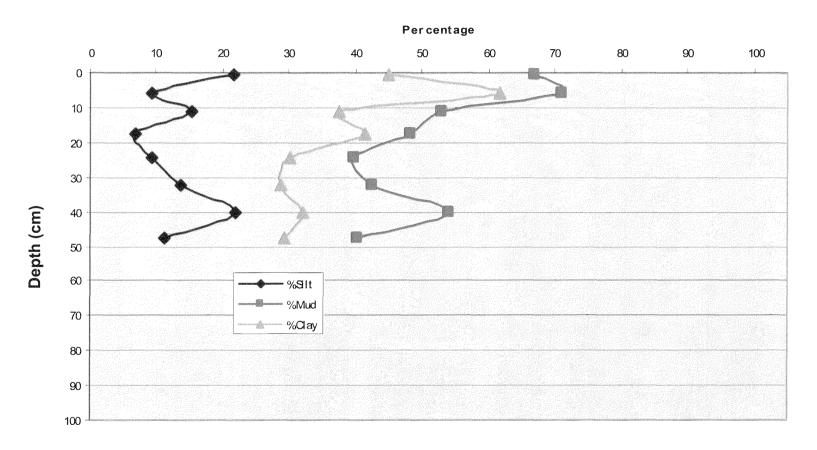


Figure 13

Brazos River - Alpha for Aquilla 3B Core

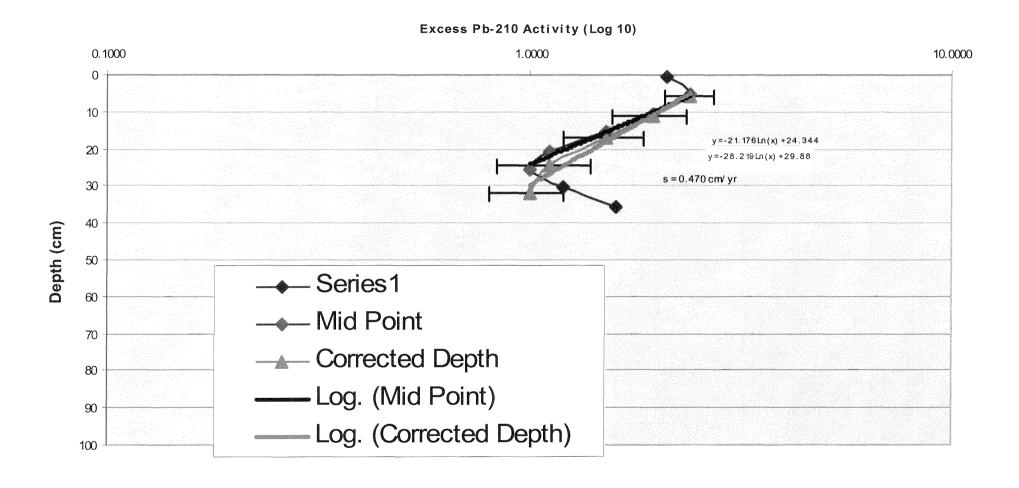
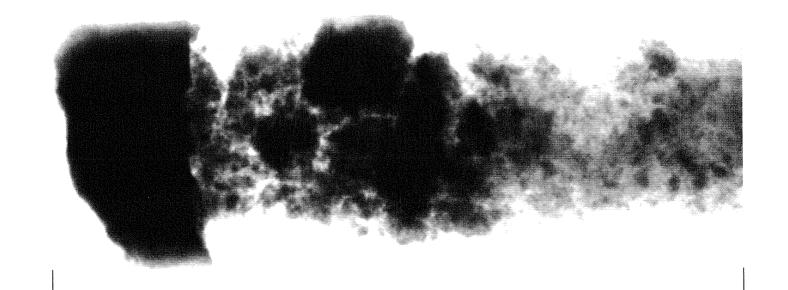


Figure 14



Lake Limestone Core 2

C C O

1 2 5 5 6 7

Grain Size Analysis of Brazos River Core LS 2B

Percentage − %SIt - %Mud %Clay Depth (cm)

Figure 17

Brazos River - Alpha for LS2B Core

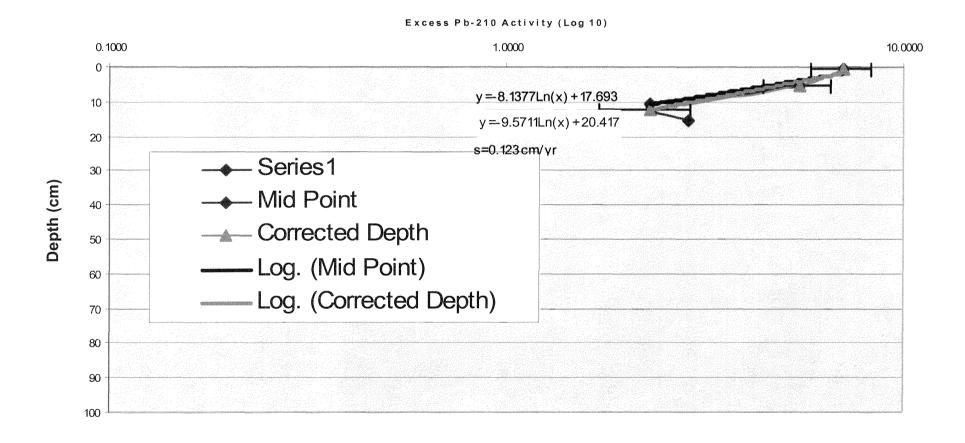
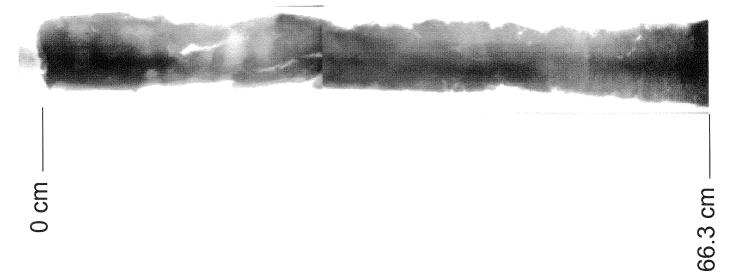


Figure 18



Lake Limestone Core 3

Water Content for LS-3 Core

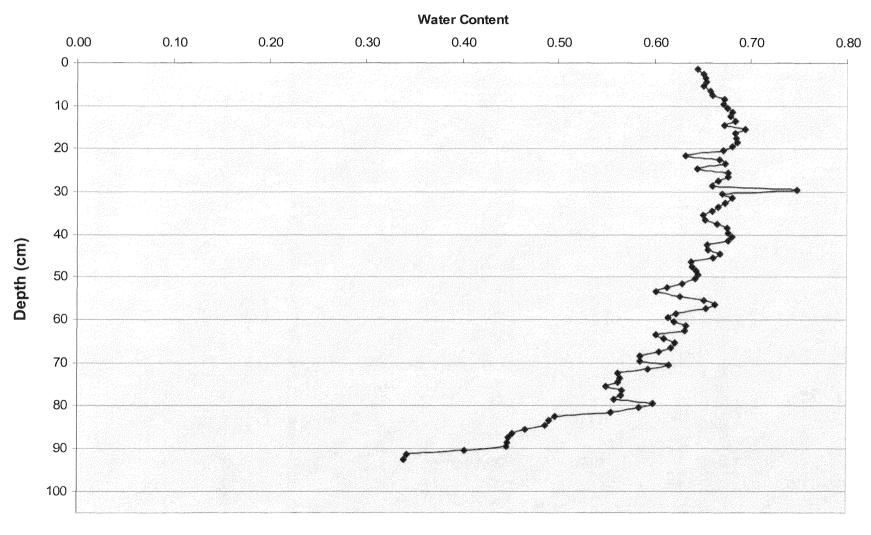


Figure 20

Grain Size Analysis of Brazos River Core LS 3

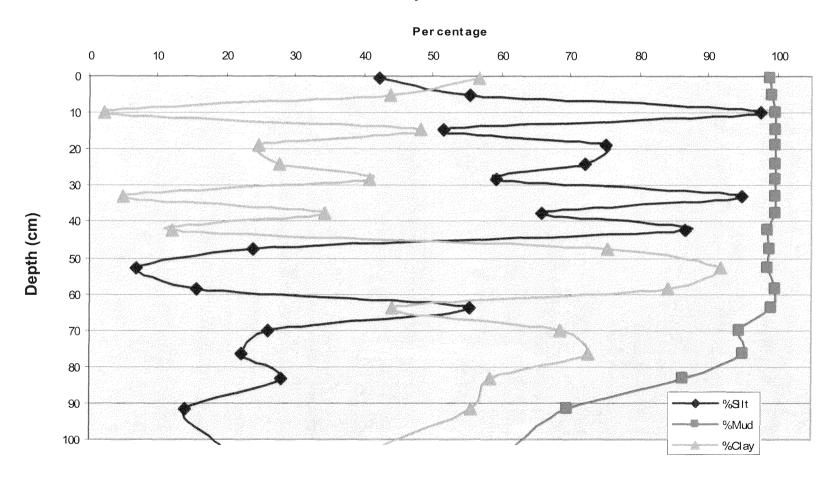


Figure 21

Brazos River - Alpha for LS-3 Core

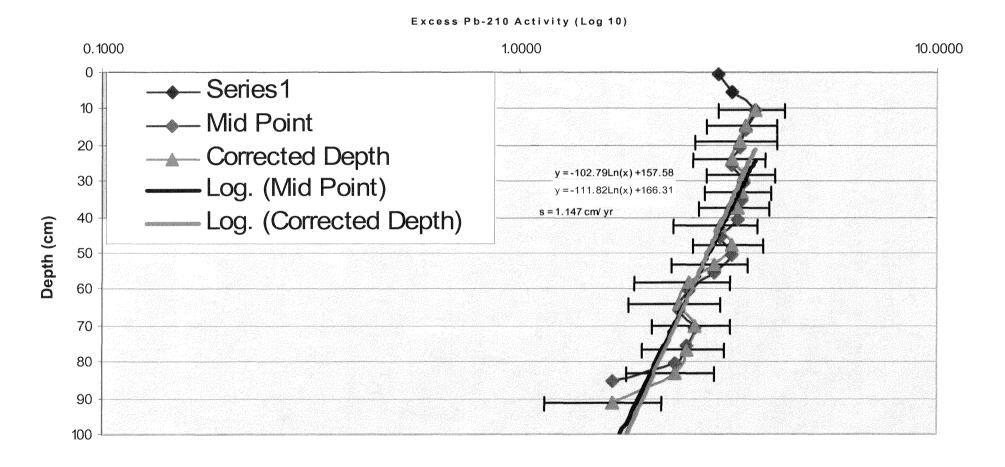
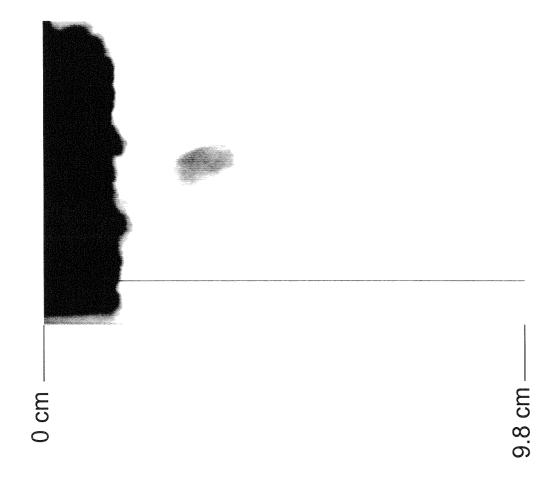


Figure 22



Lake Limestone Core 4

Figure 24

Grain Size Analysis of Brazos River Core LS 4

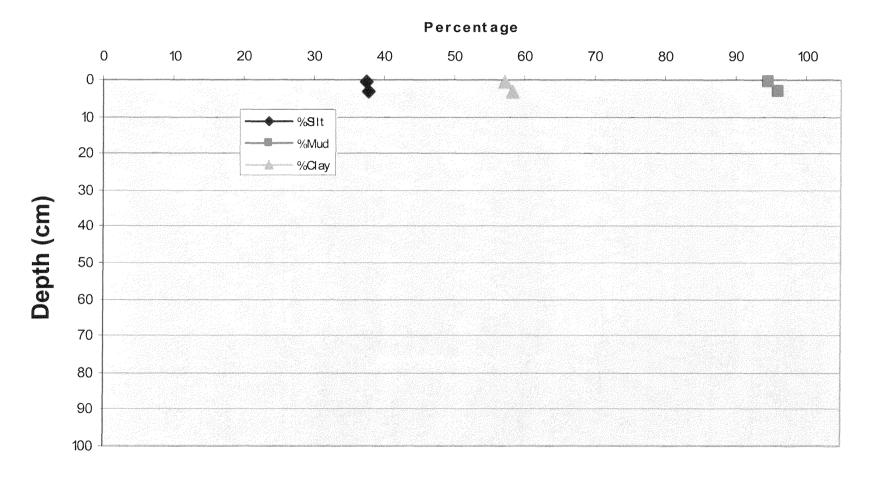
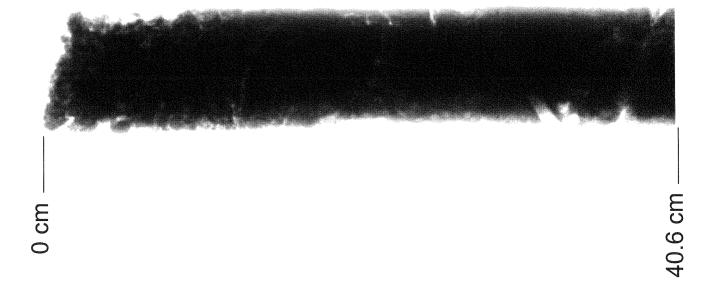


Figure 25



Lake Proctor Core 1B

Figure 27

Grain Size Analysis of Brazos River Core Proctor 1B

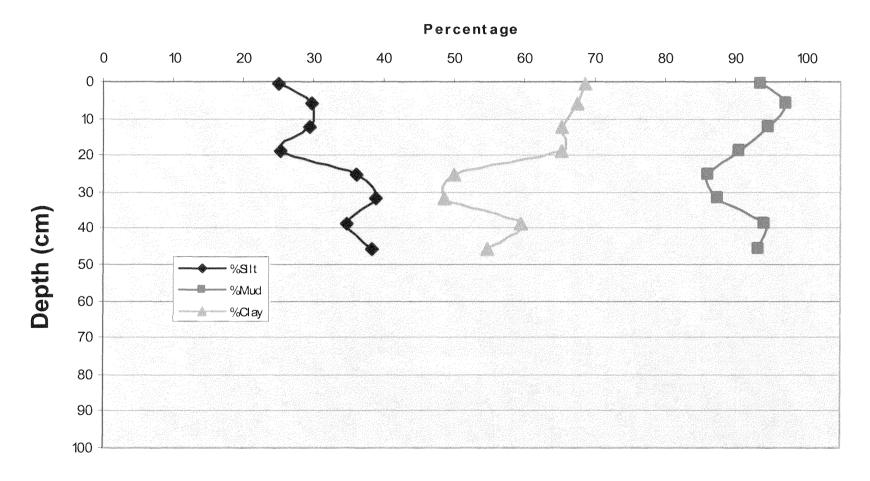


Figure 28

Brazos River - Alpha for Proctor 1B

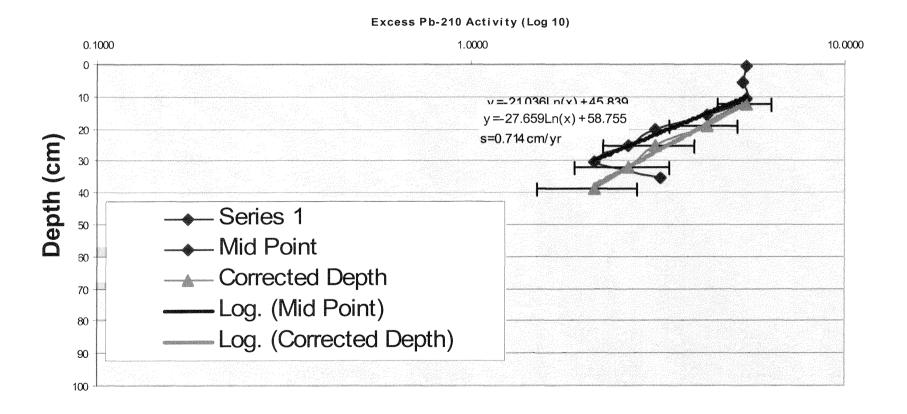


Figure 29

22.5 cm

Granger Lake Core 1B

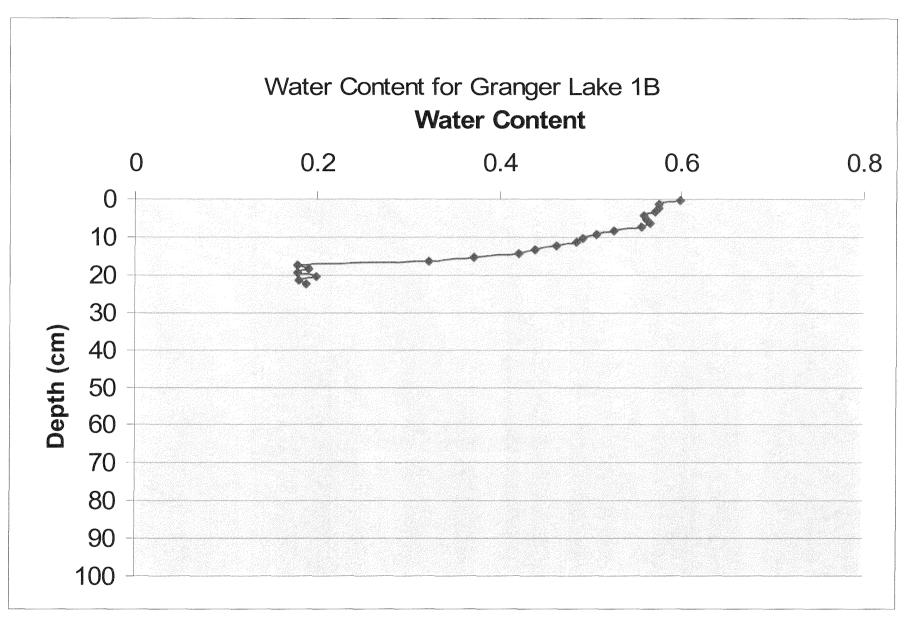


Figure 31

Grain Size Analysis of Brazos River Core Granger 1B

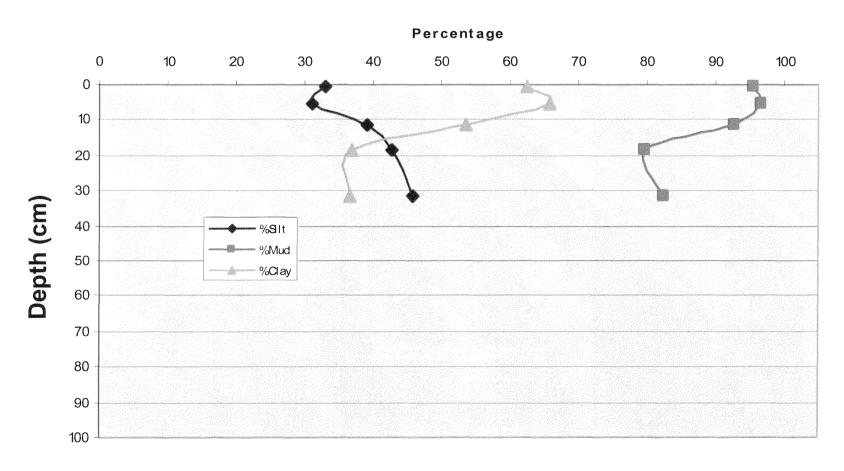


Figure 32

Brazos River - Alpha for Granger 1B Core

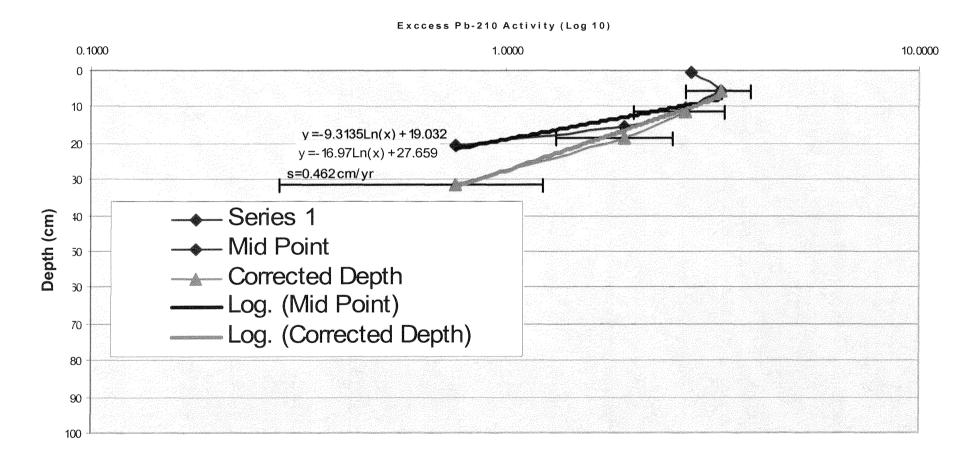
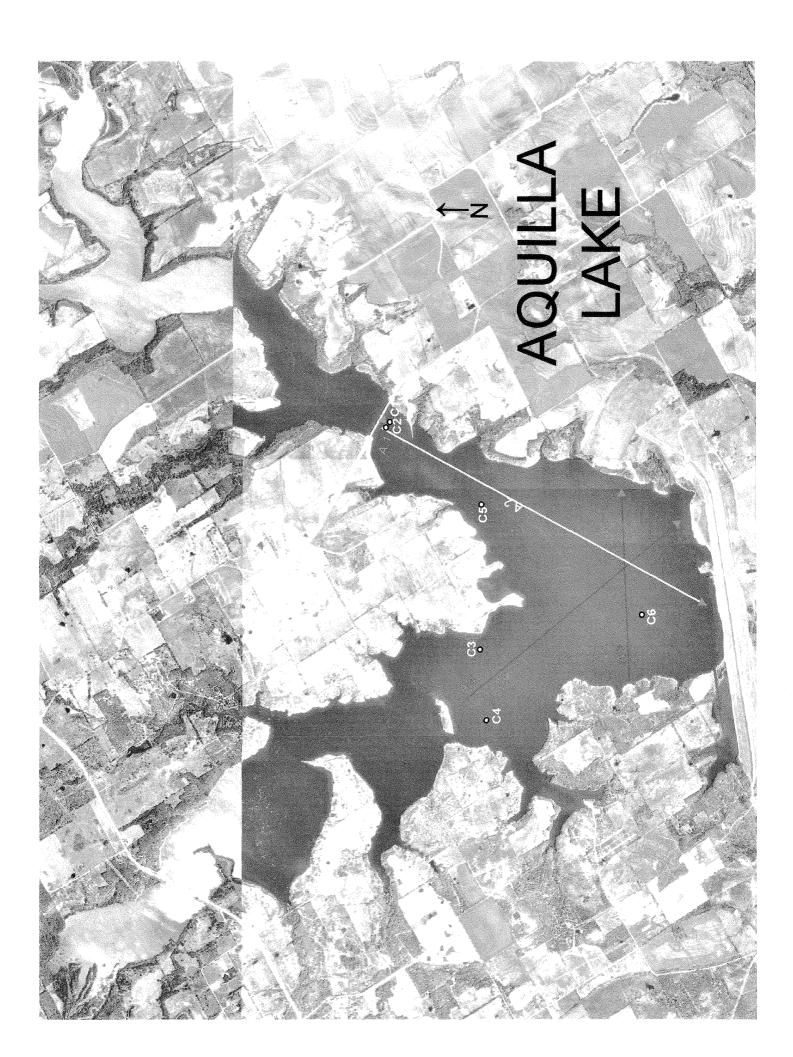
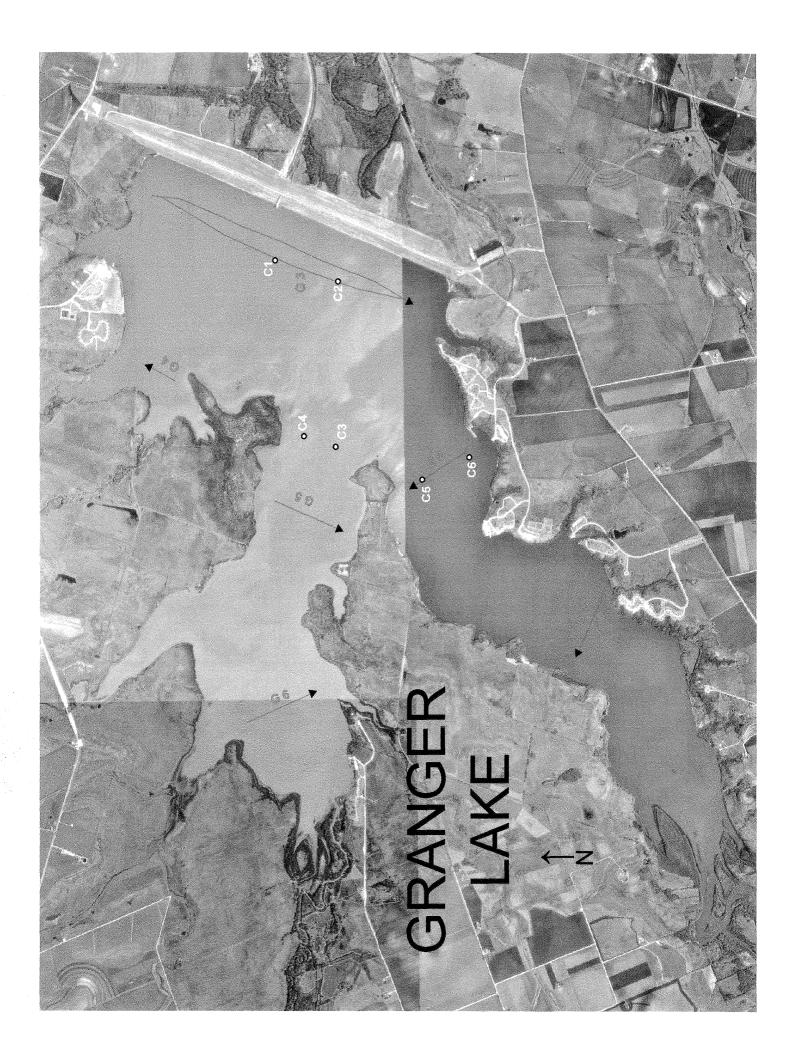
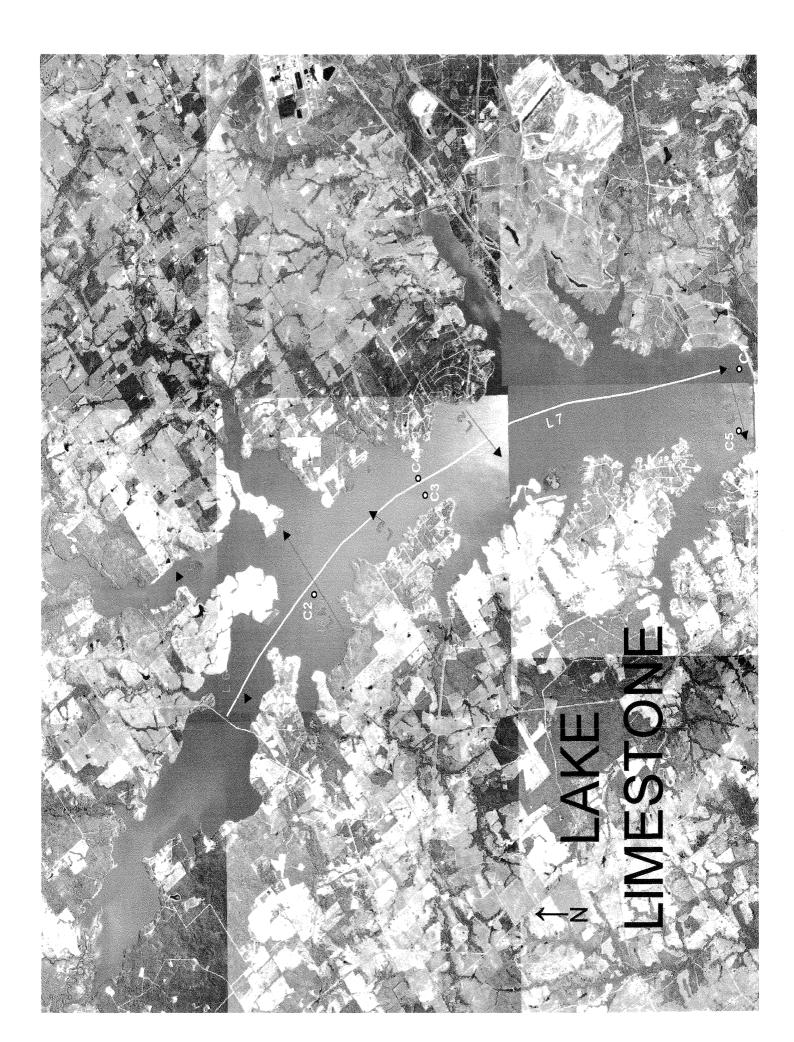
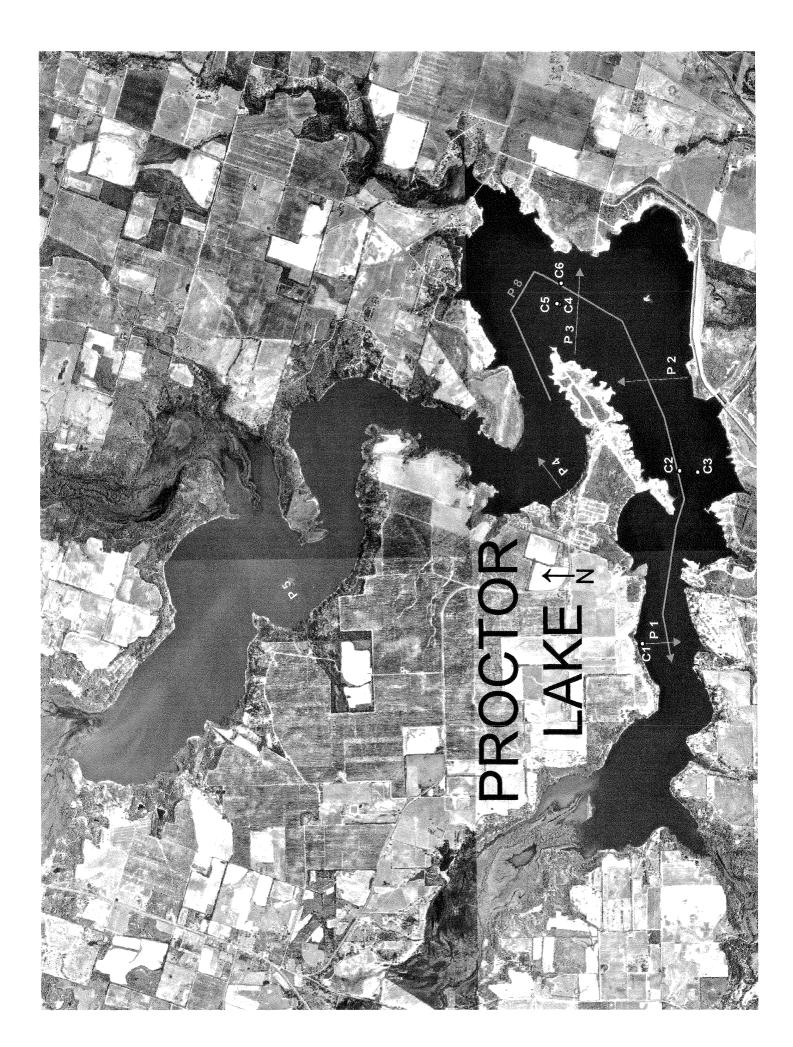


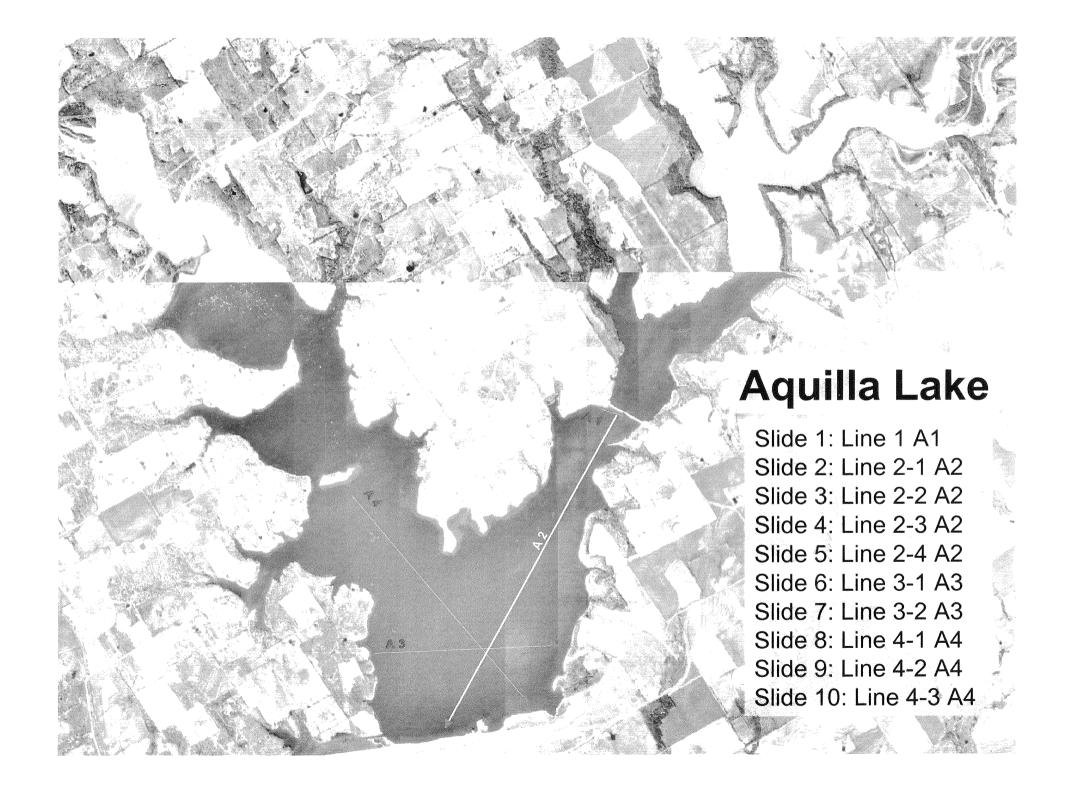
Figure 33

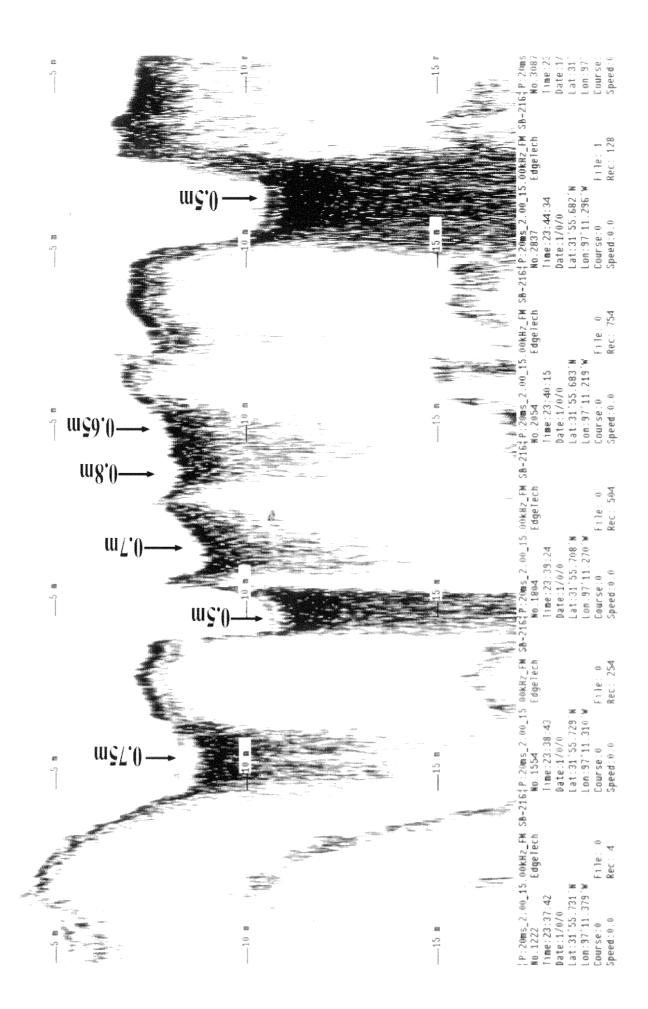


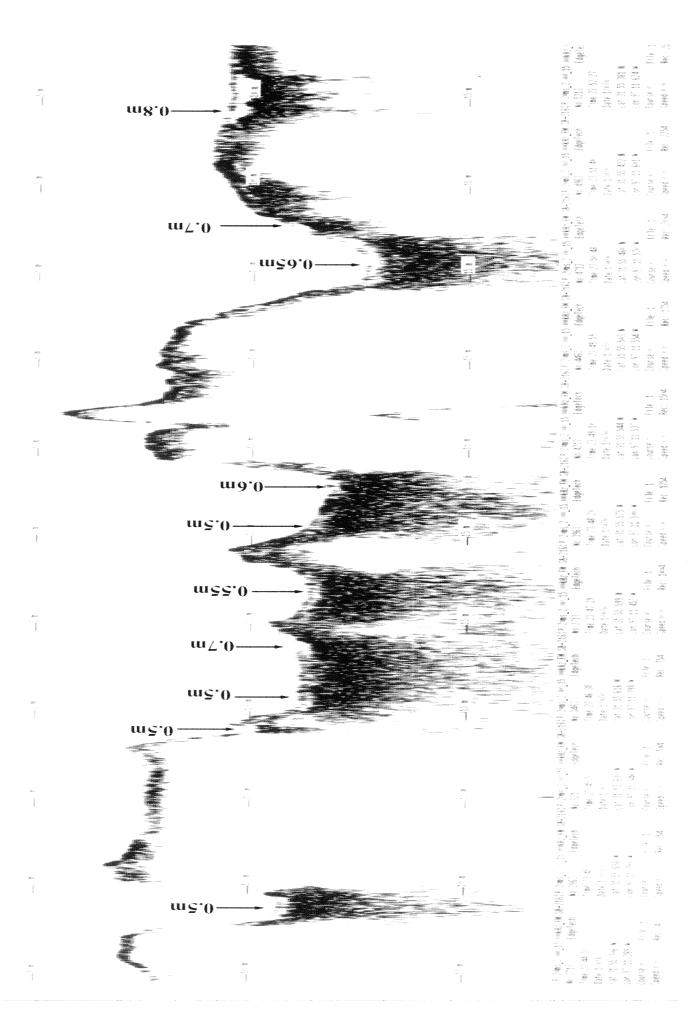


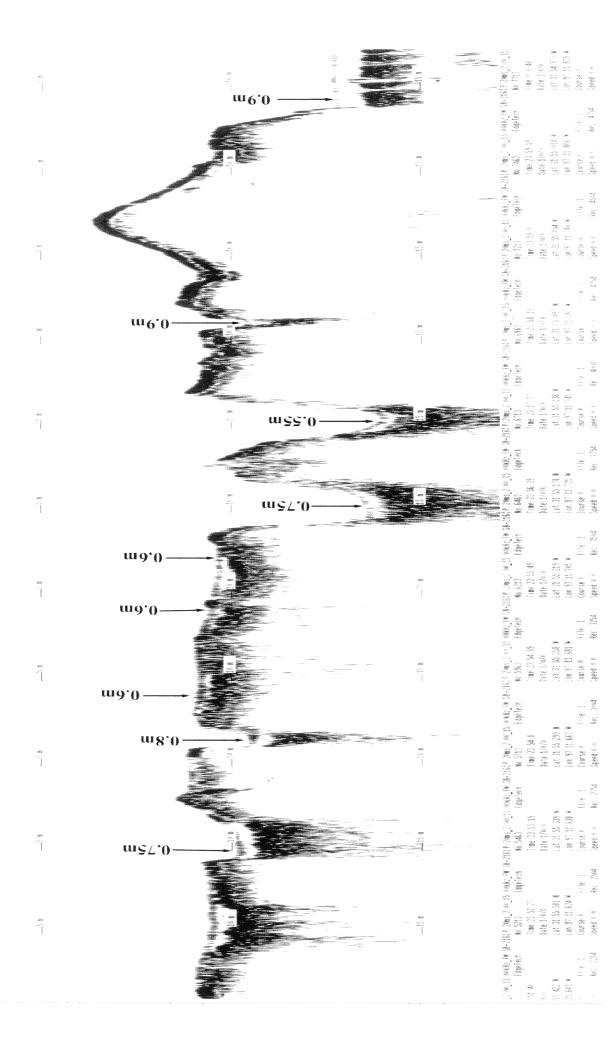


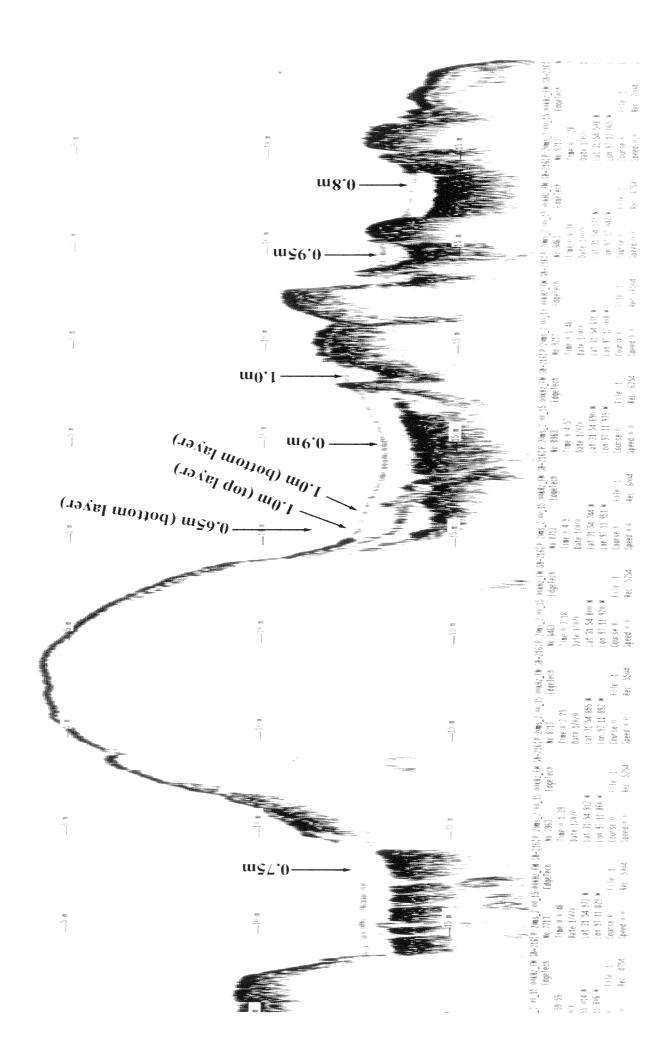




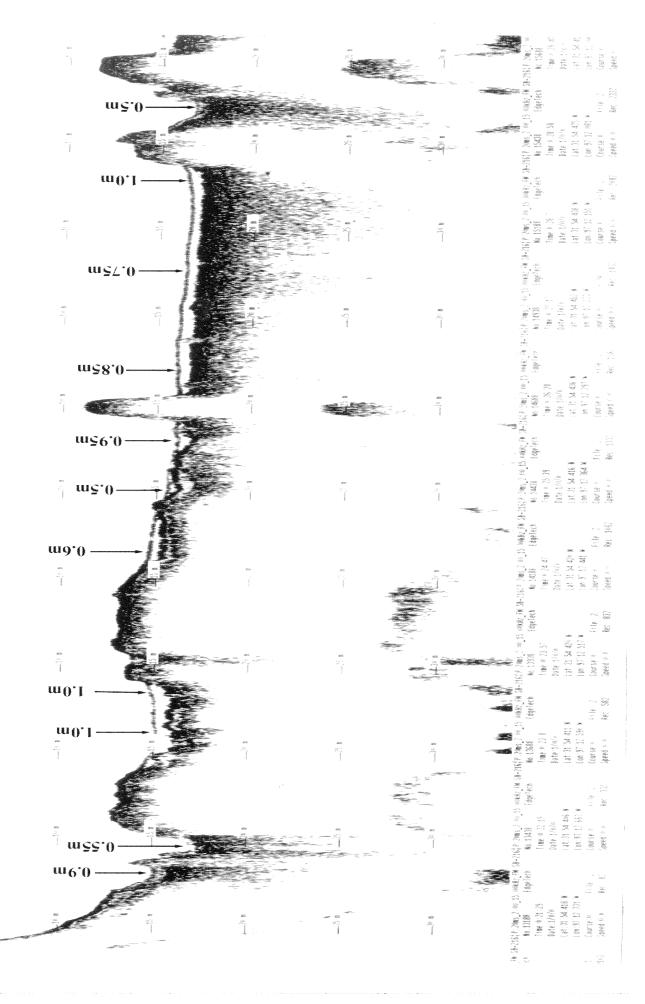




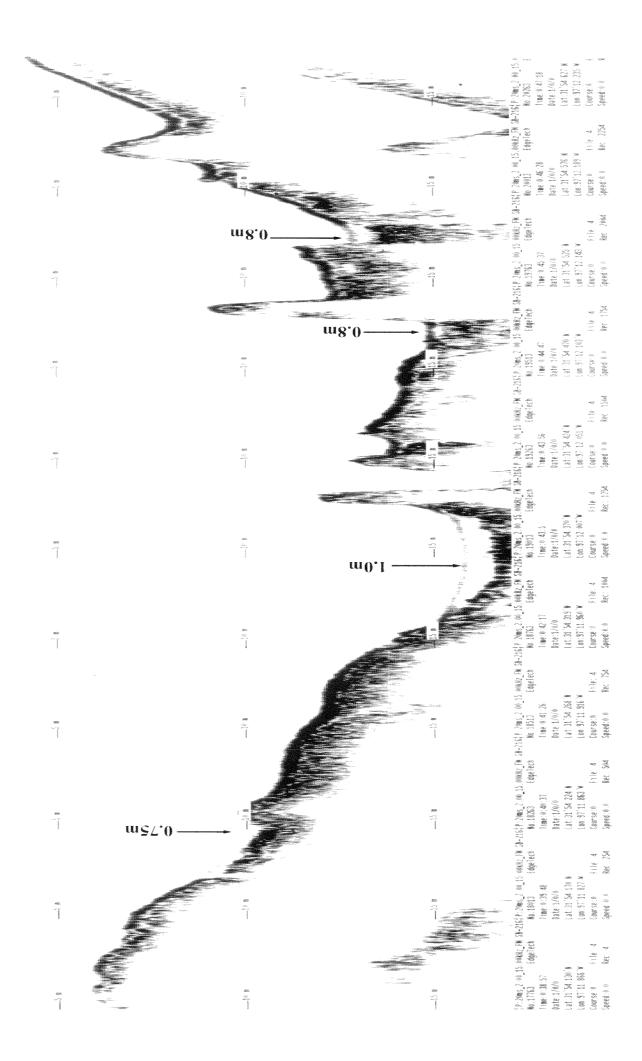


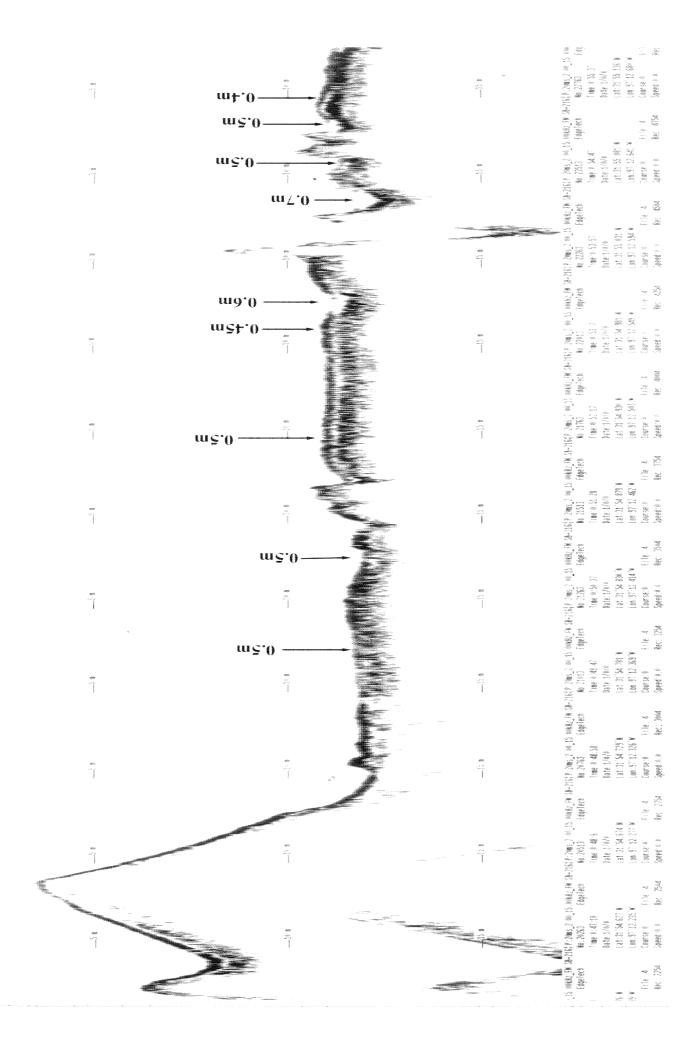


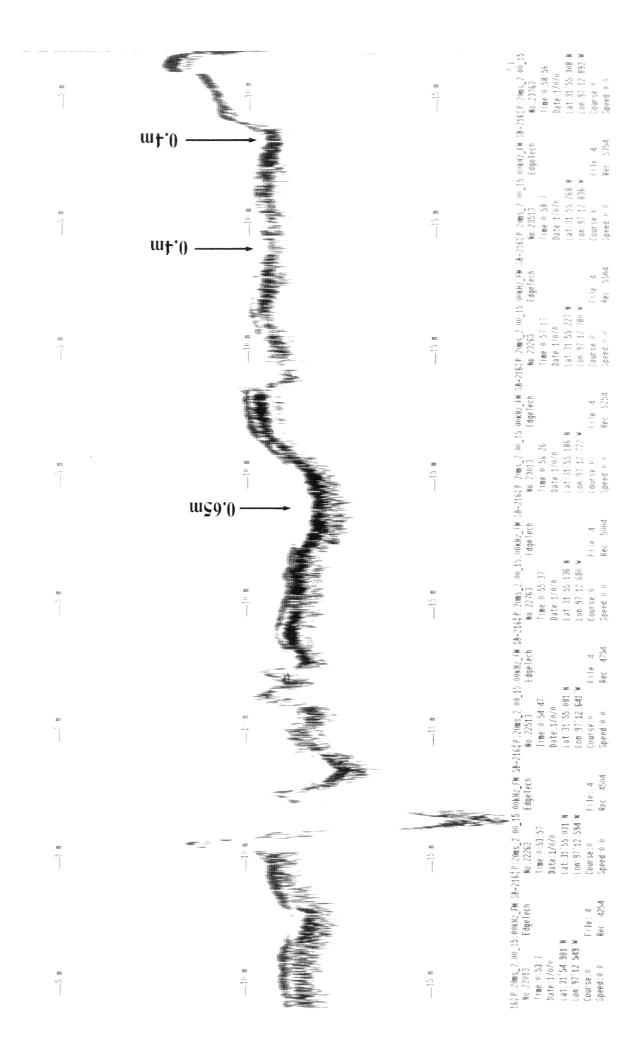
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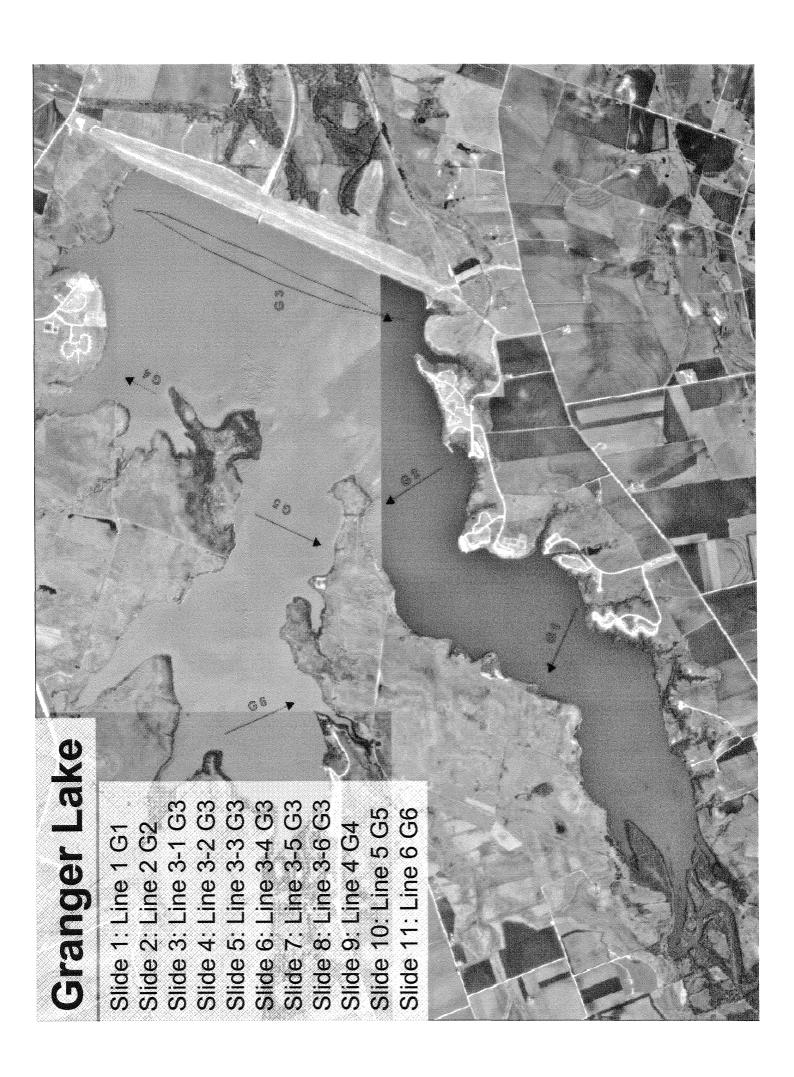


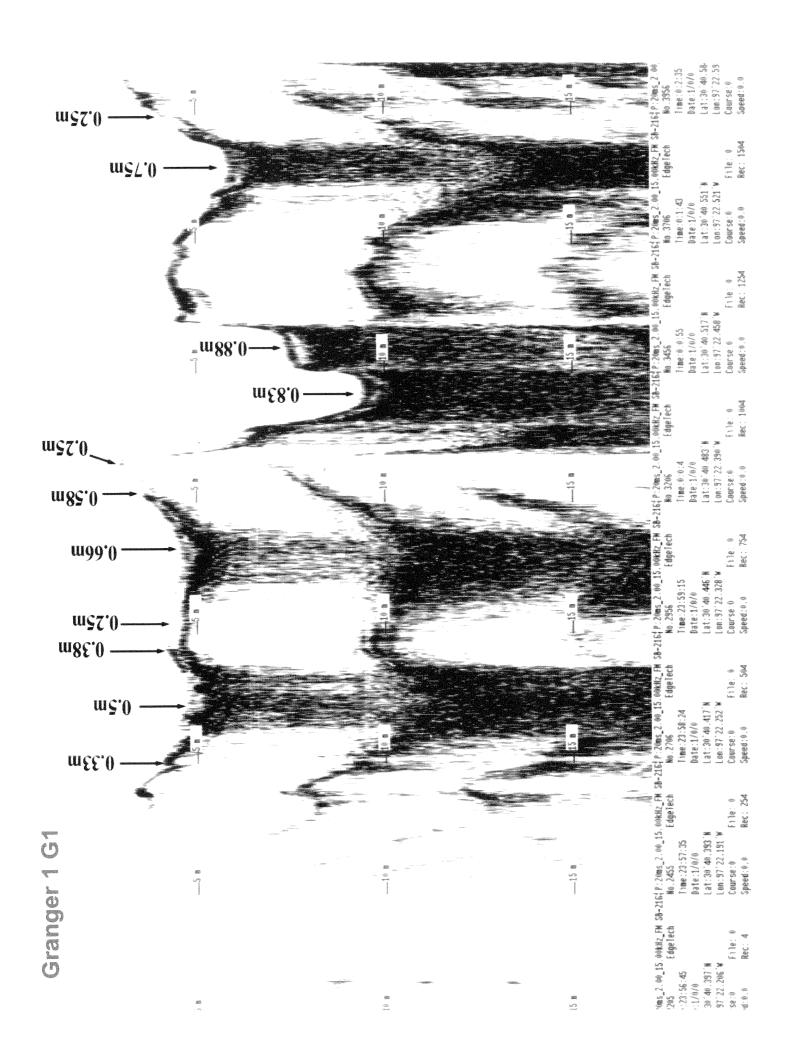
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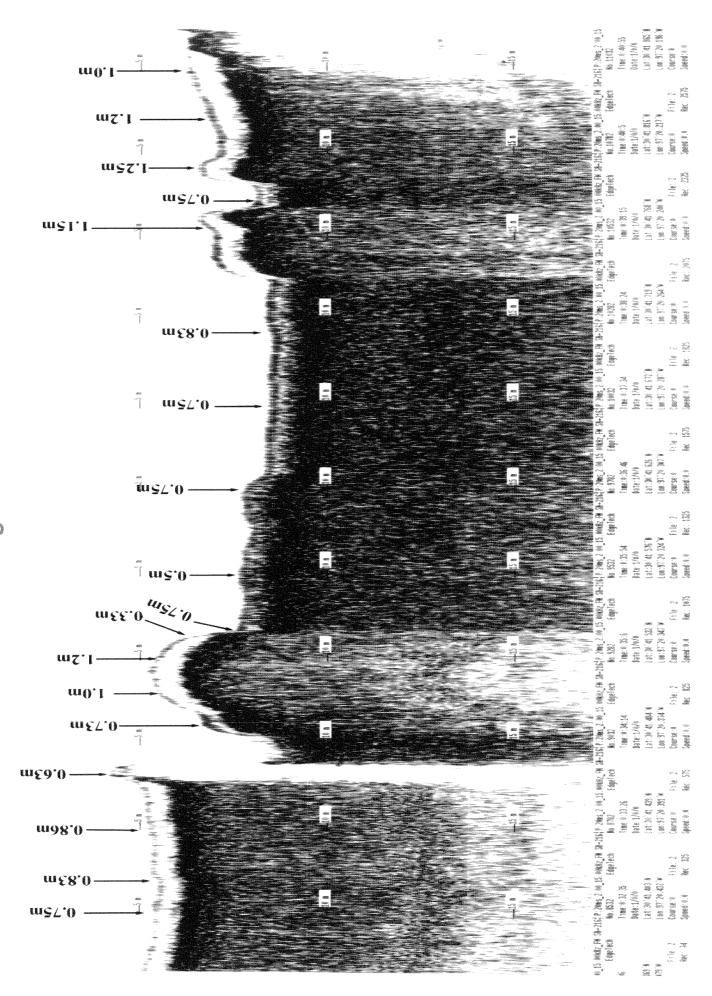


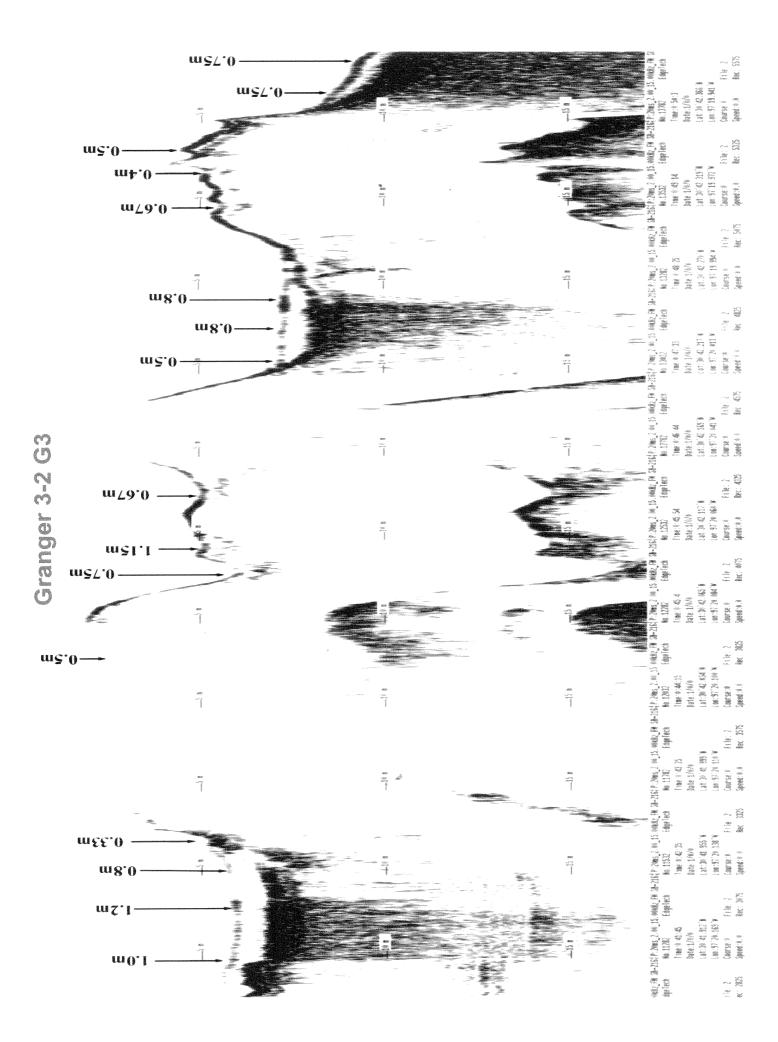


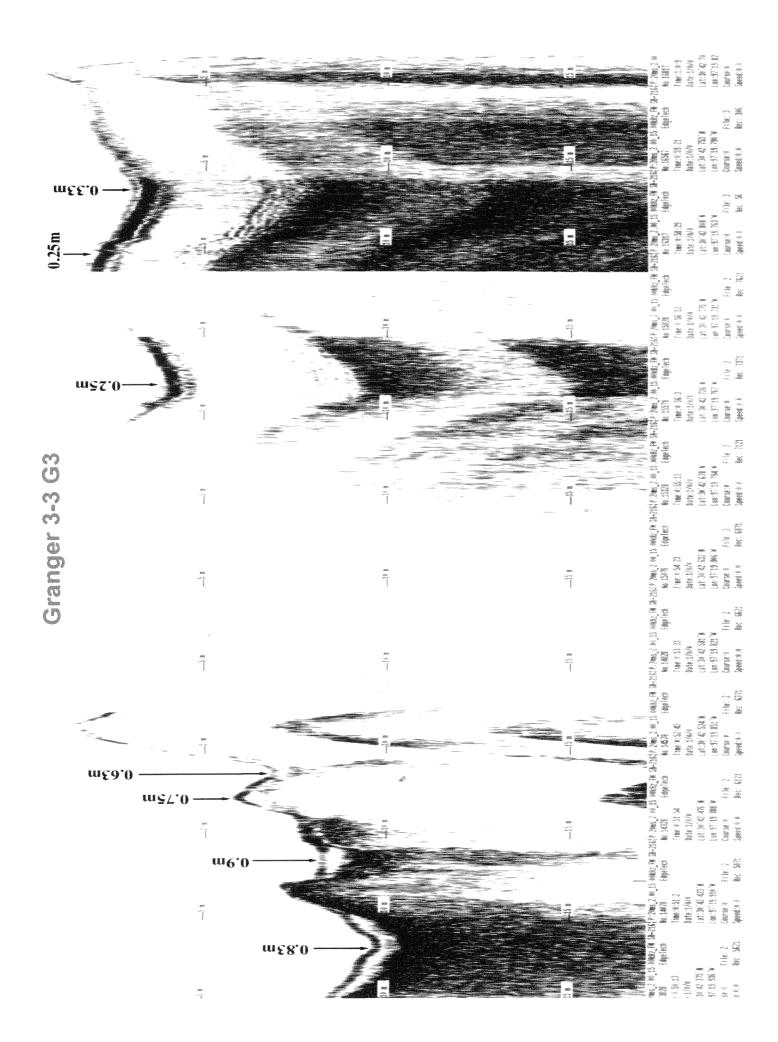


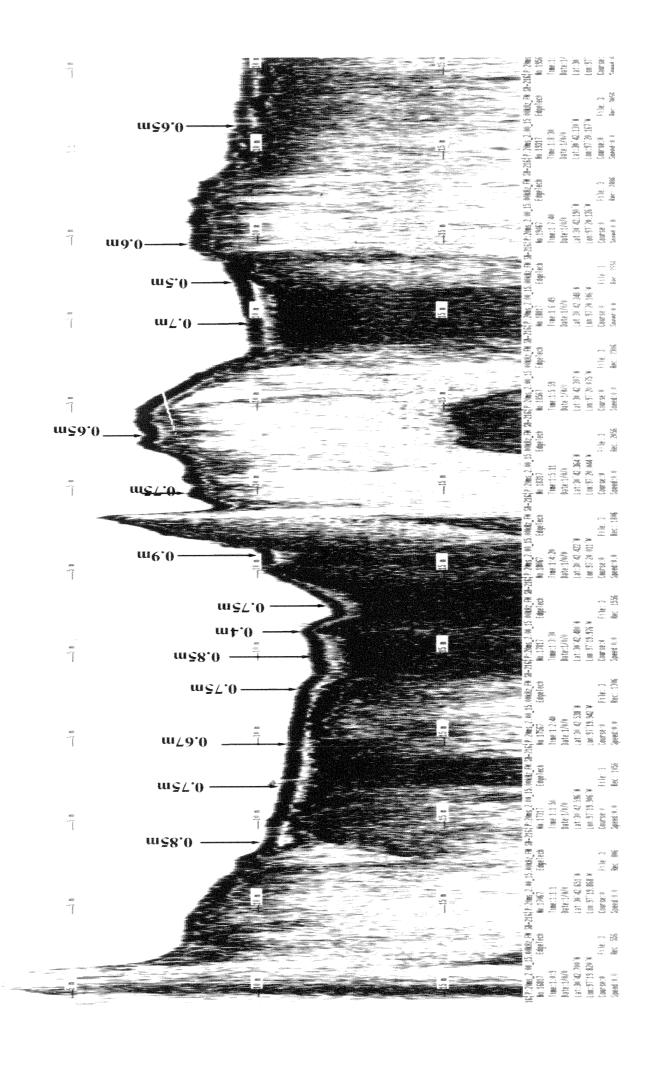


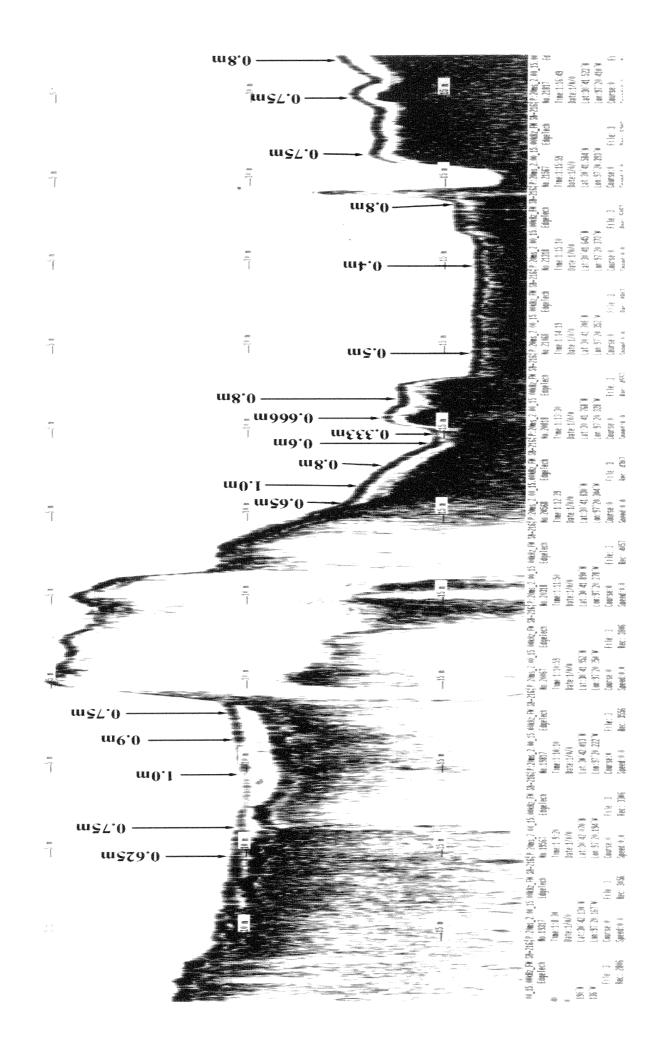


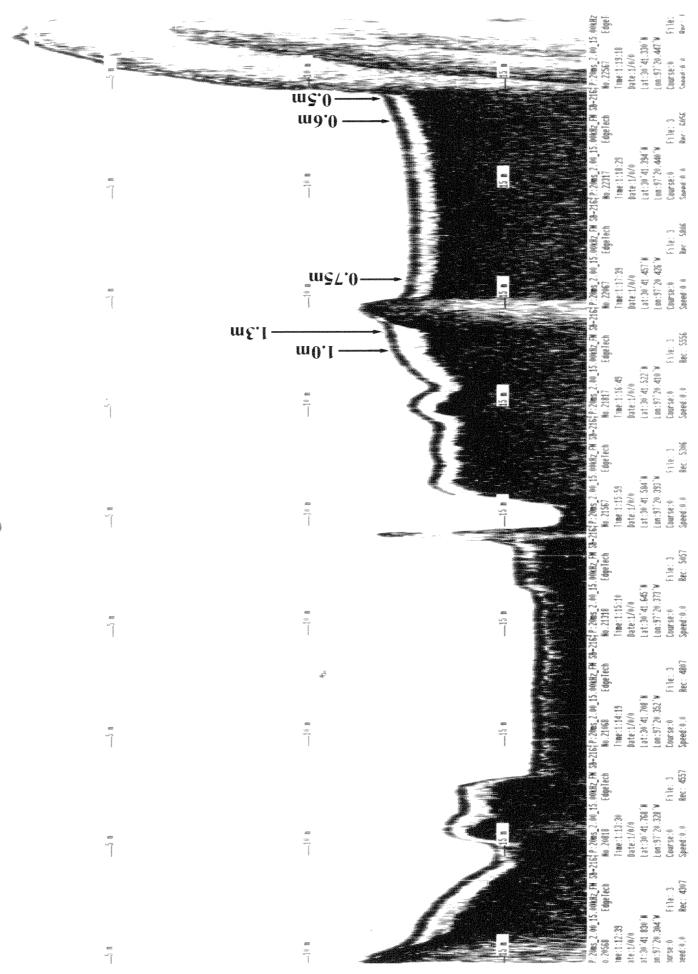


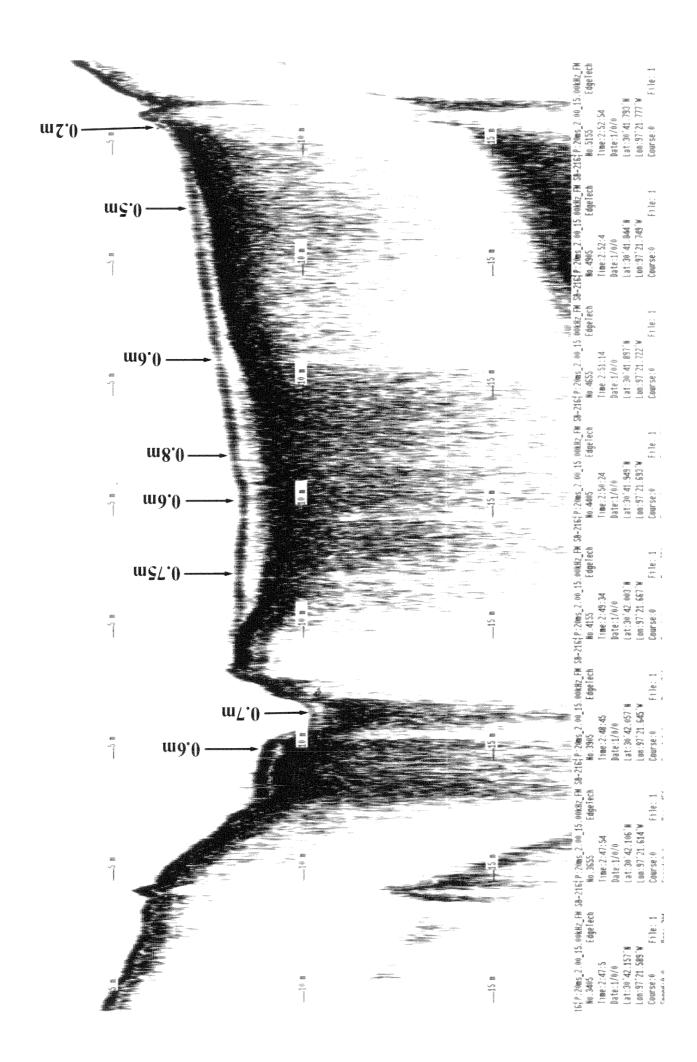




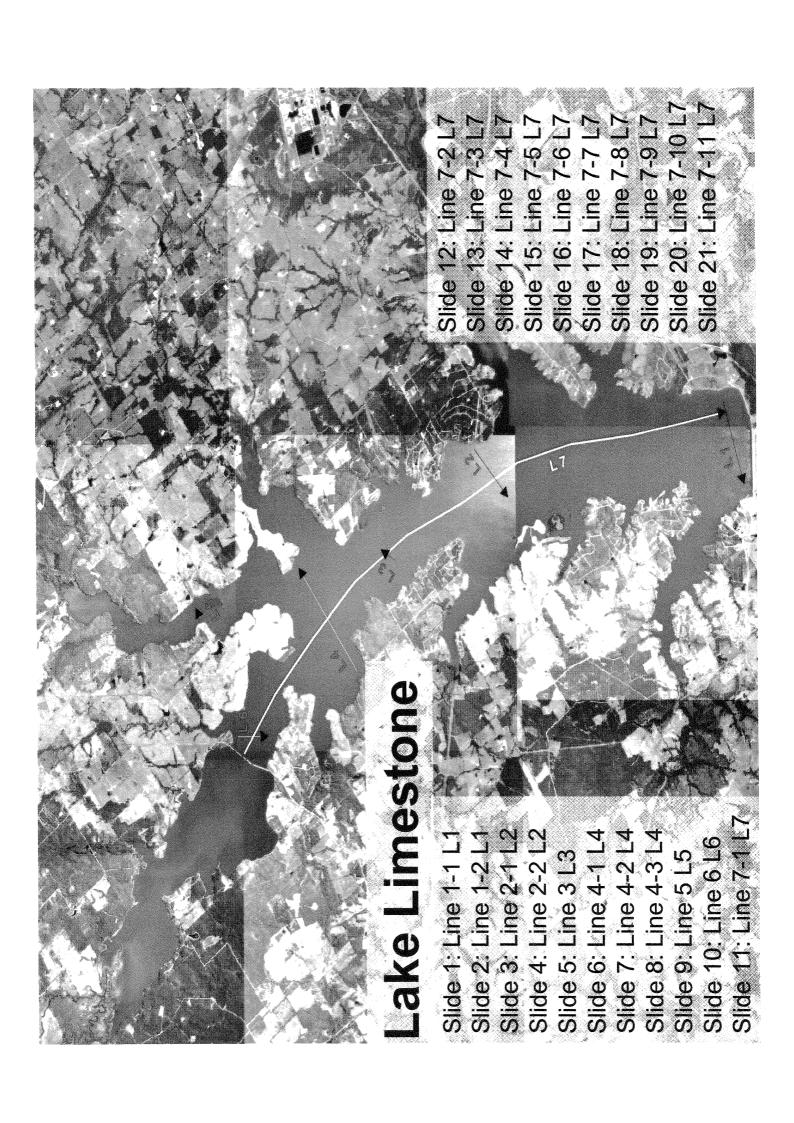








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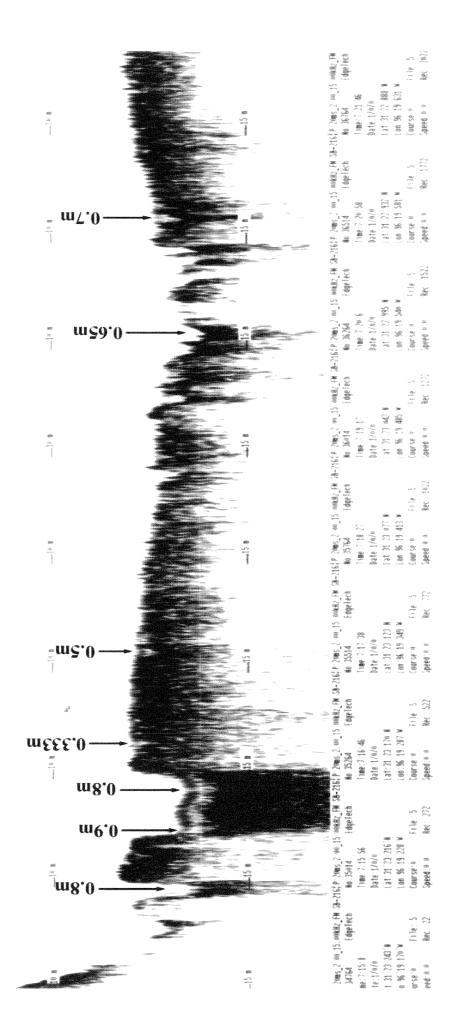


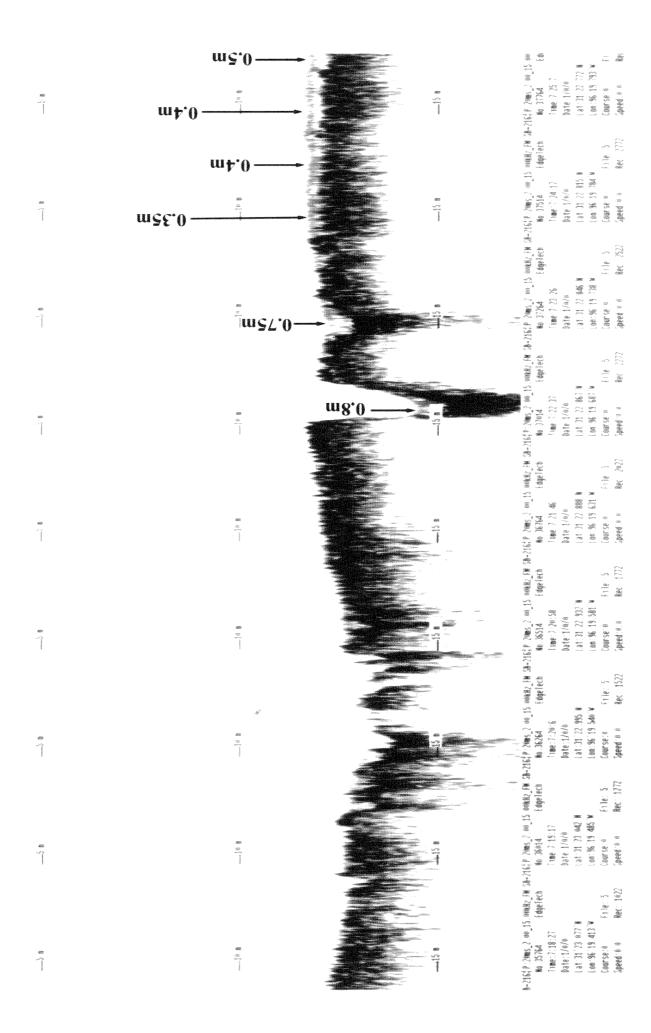
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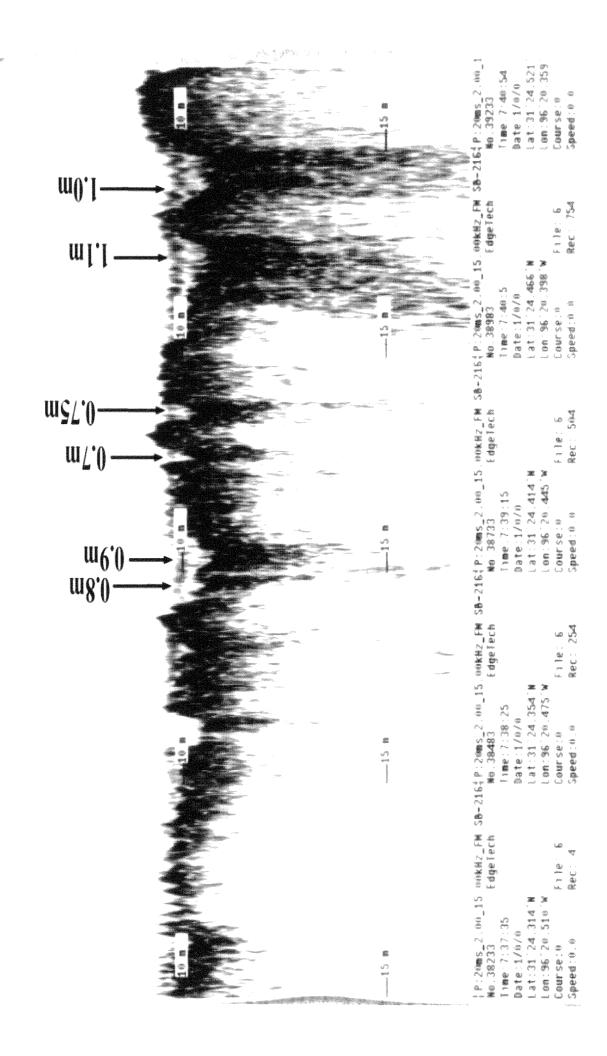
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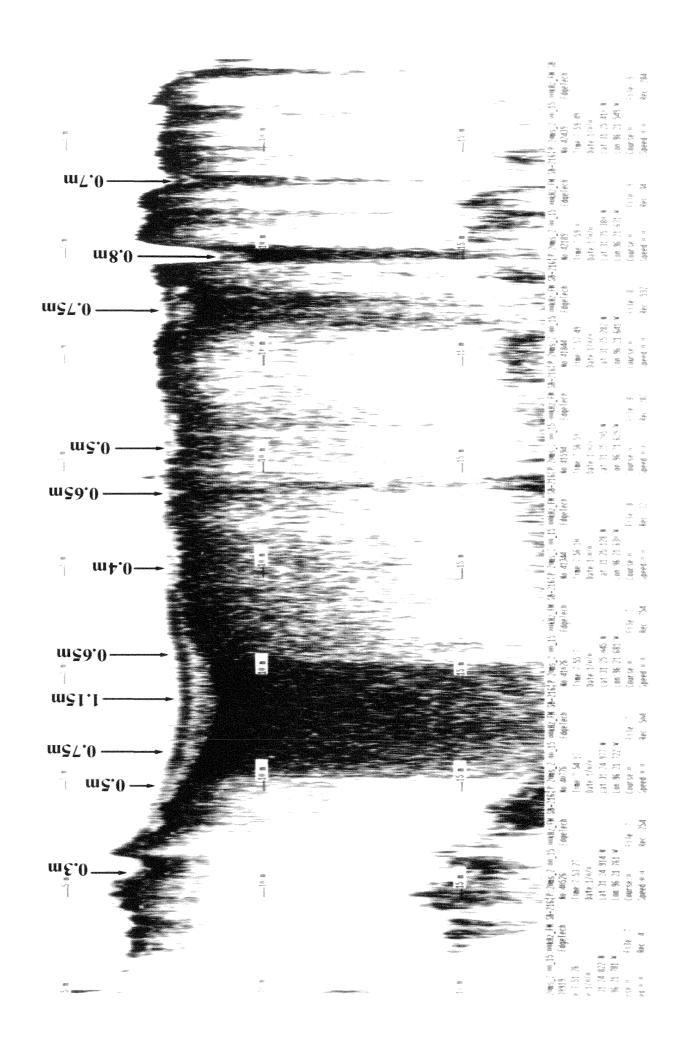
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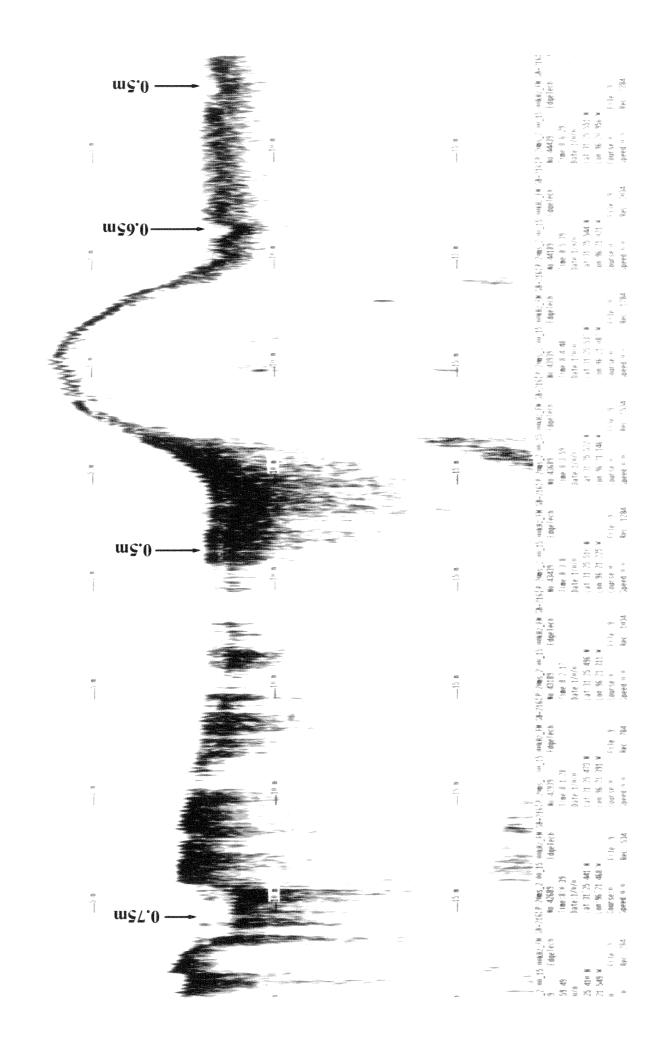
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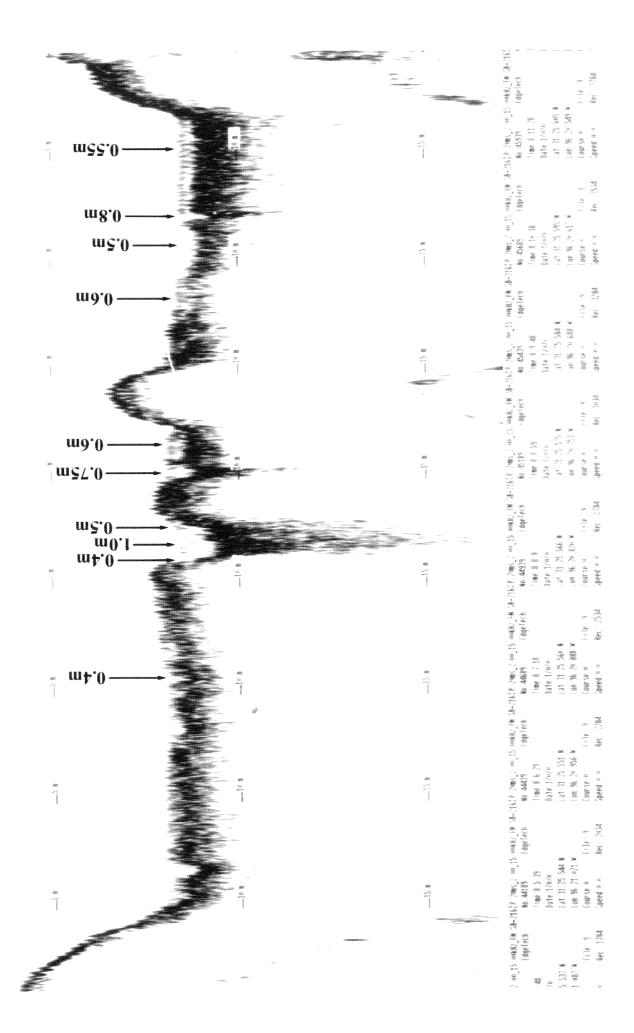


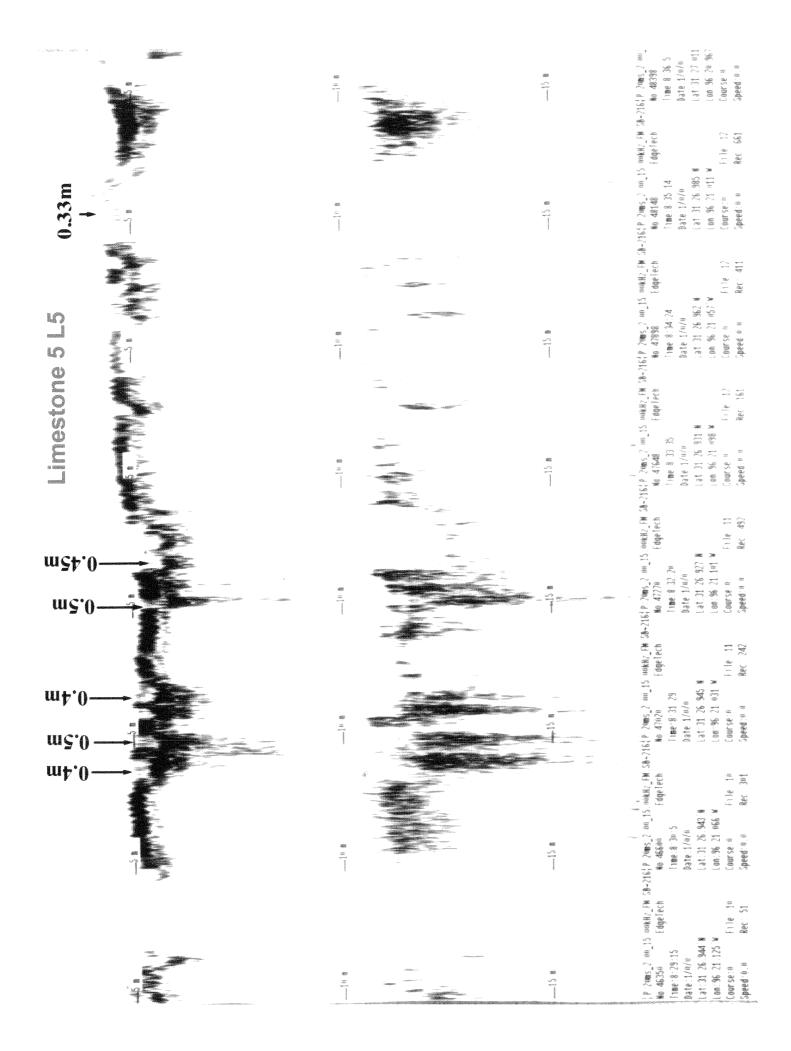


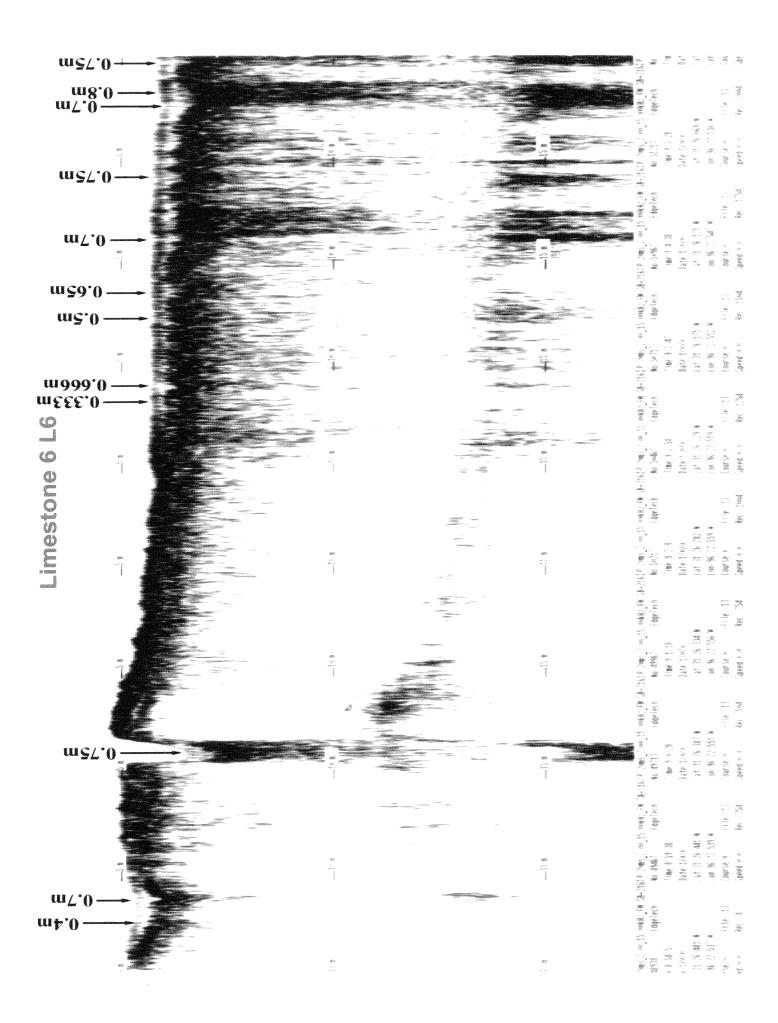




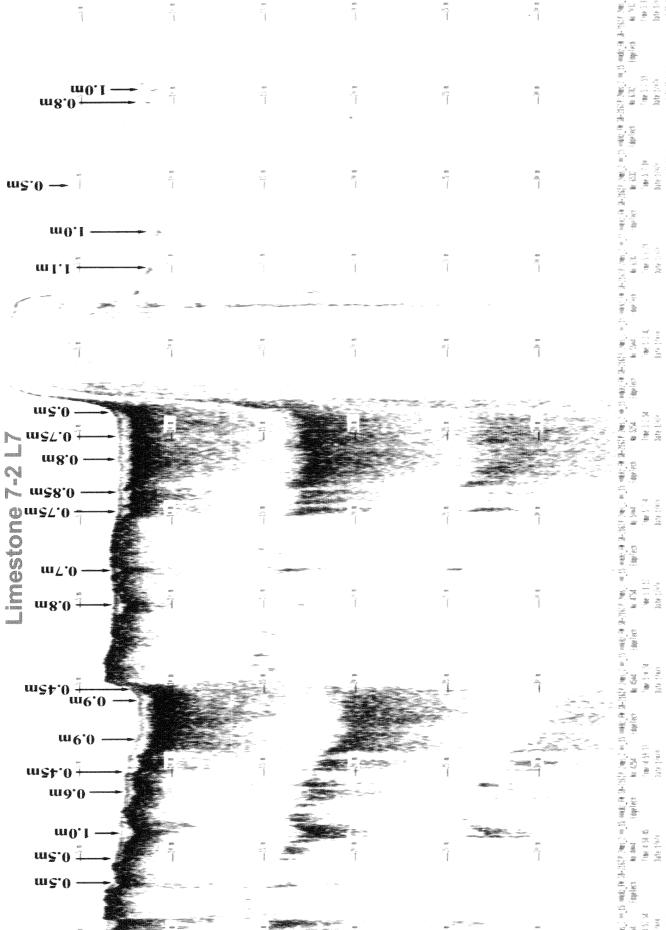






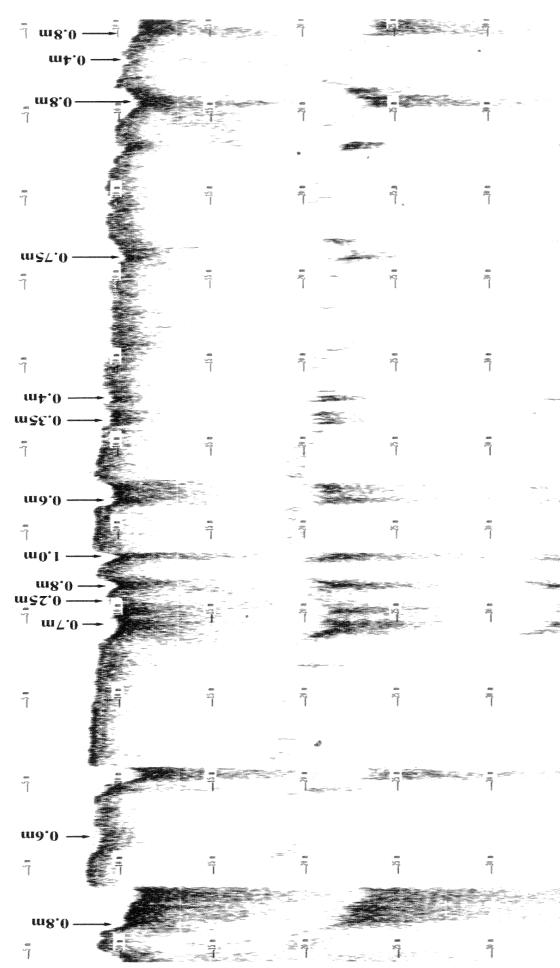


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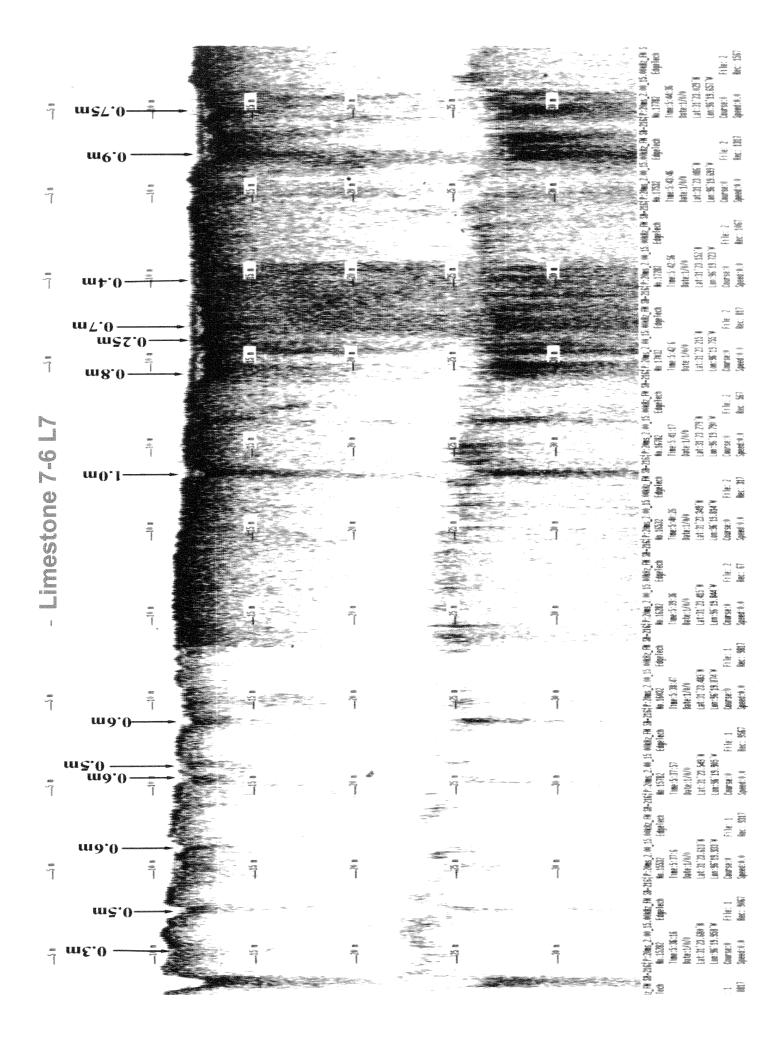


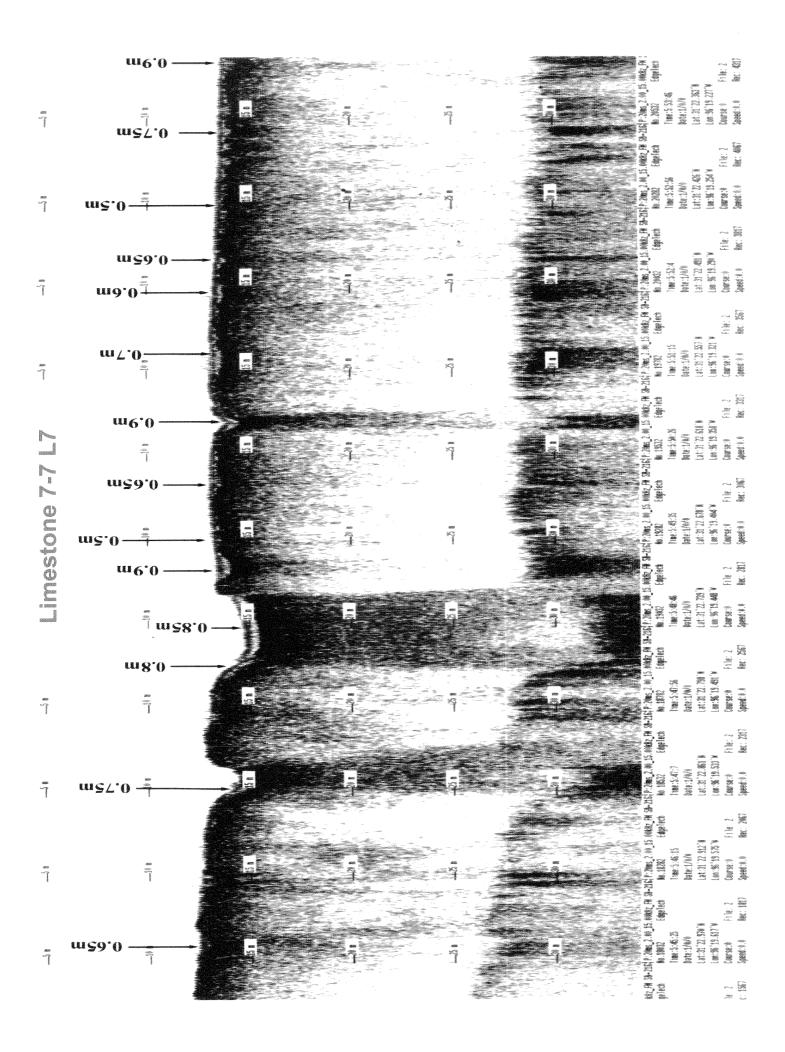
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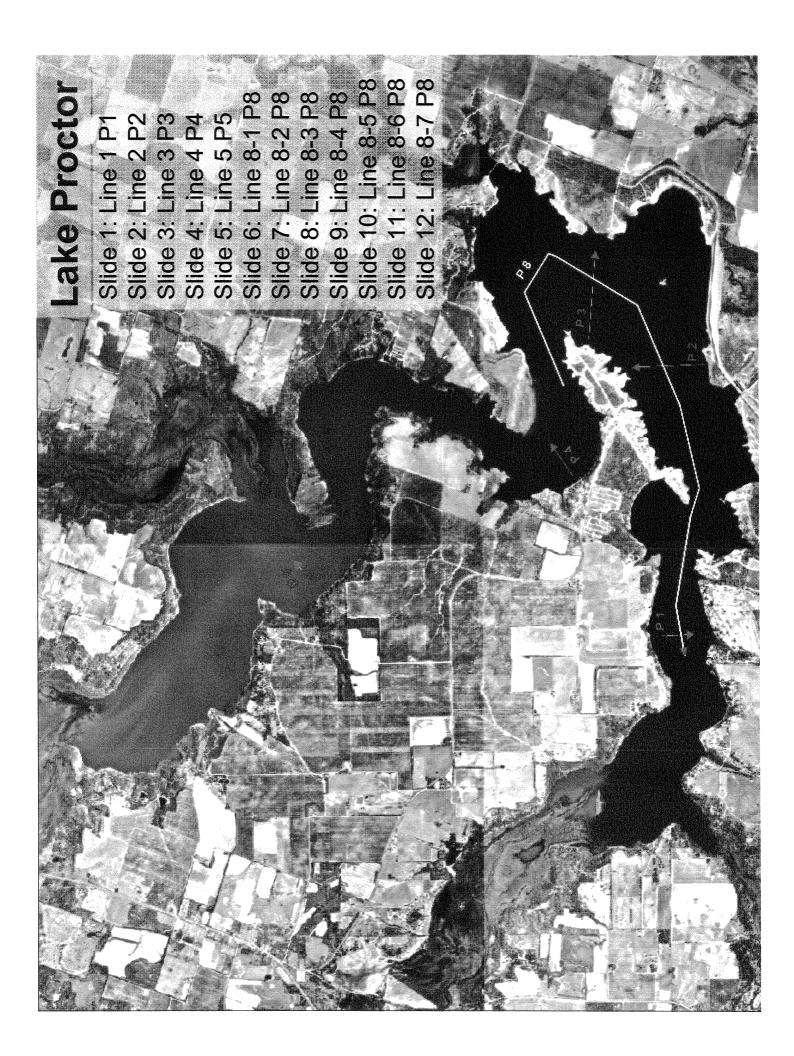
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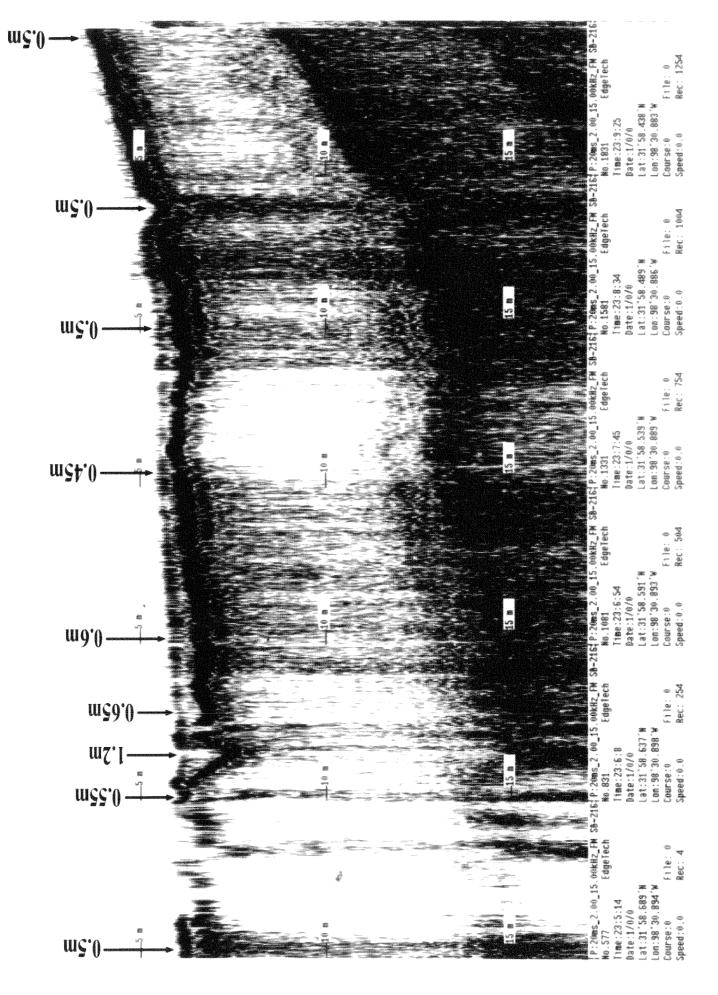


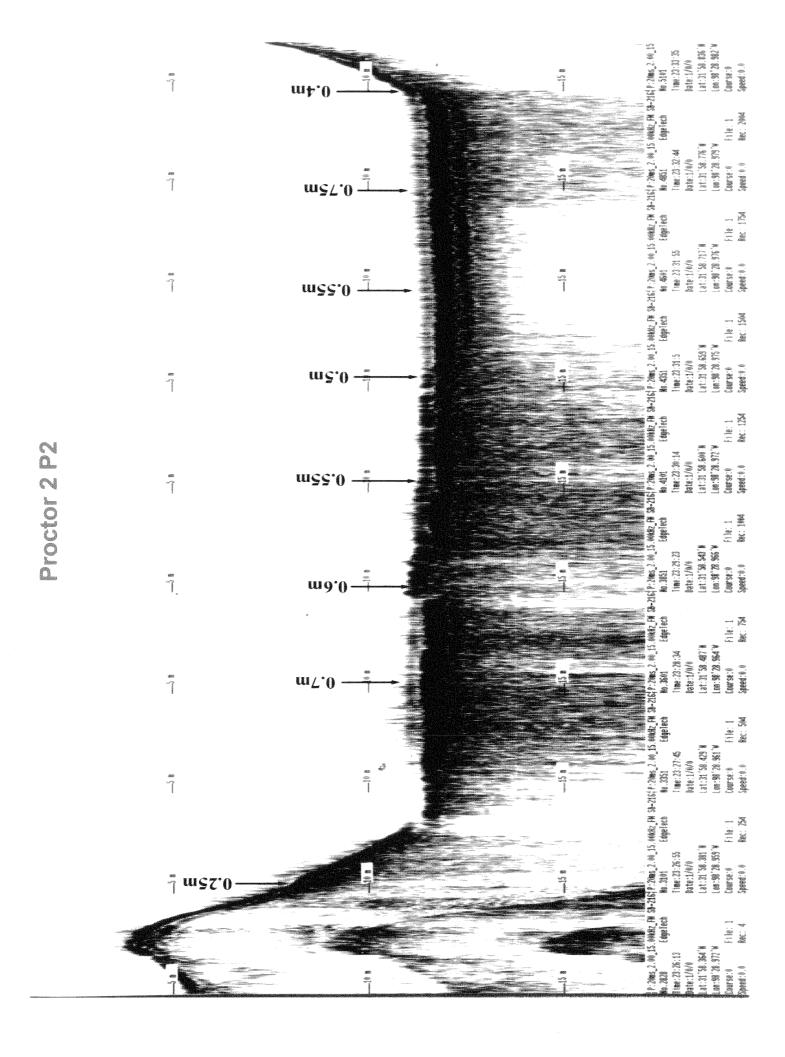


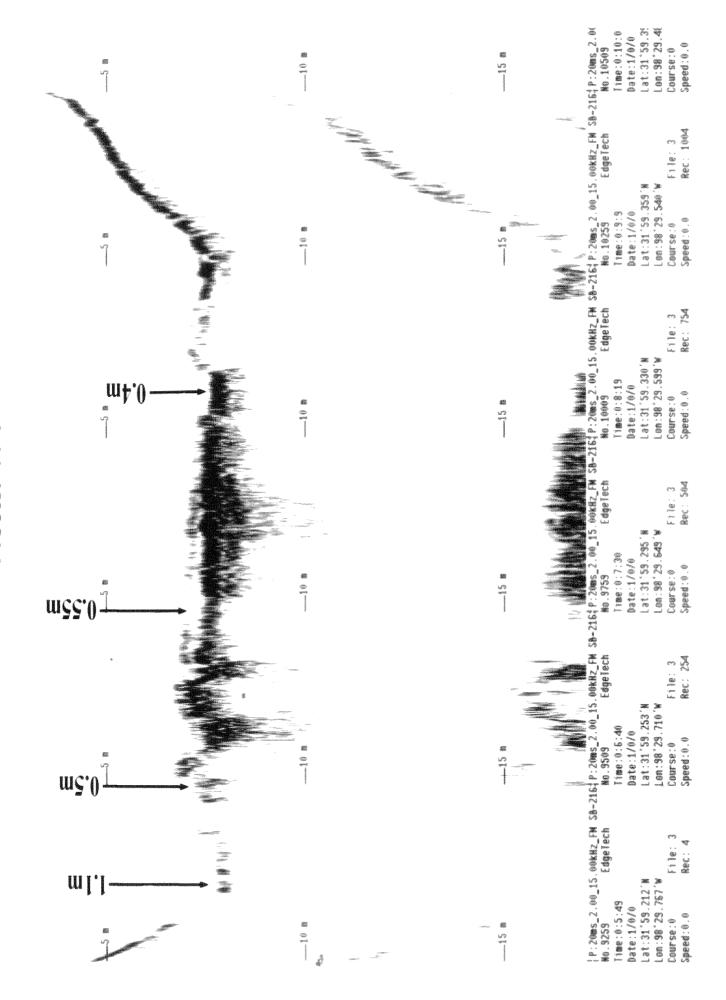
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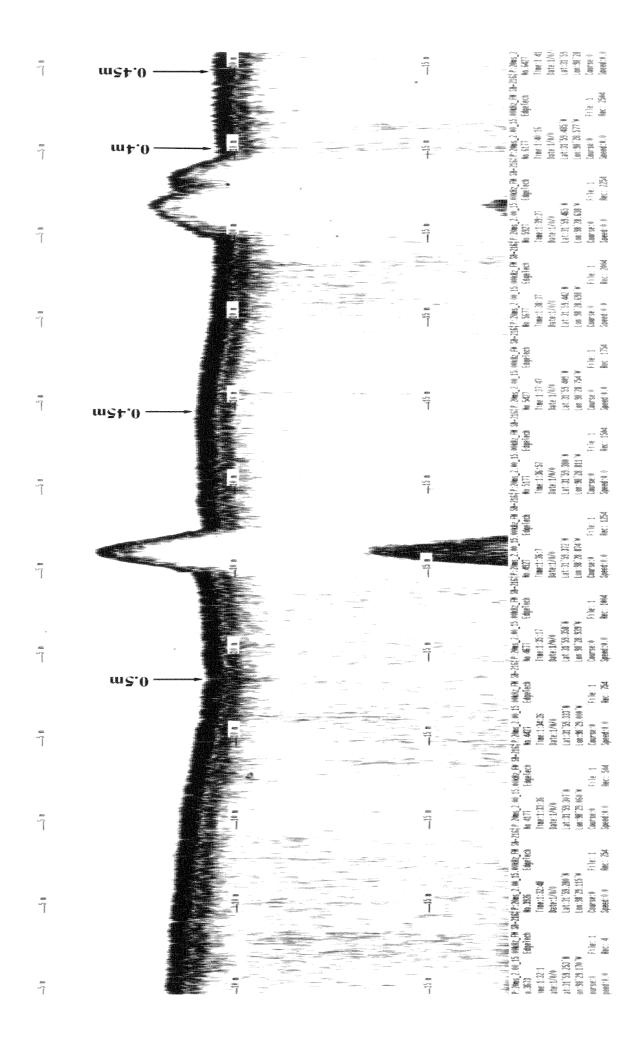


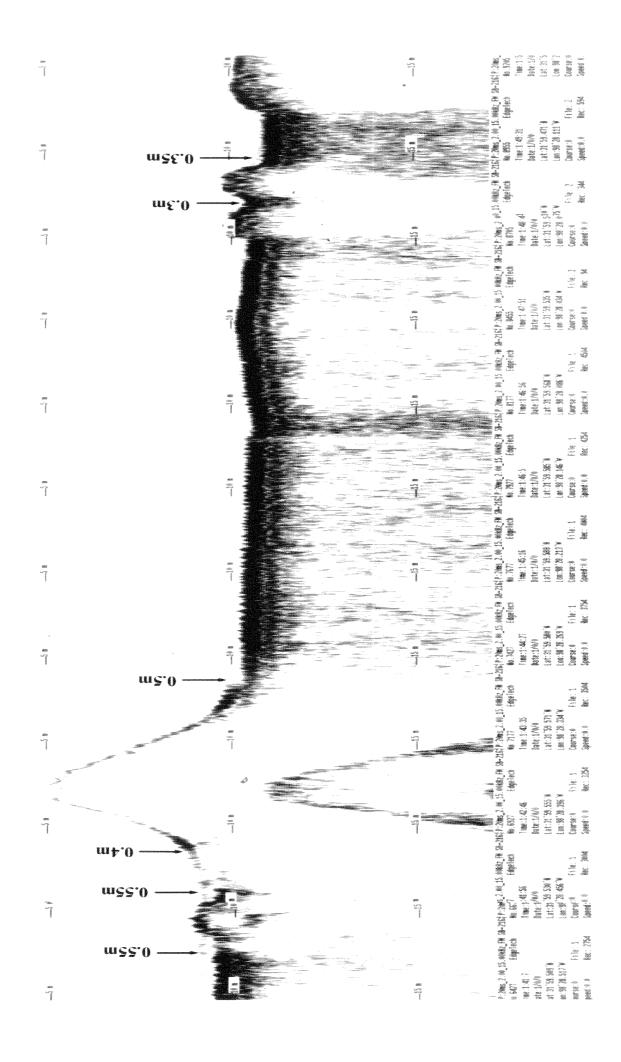


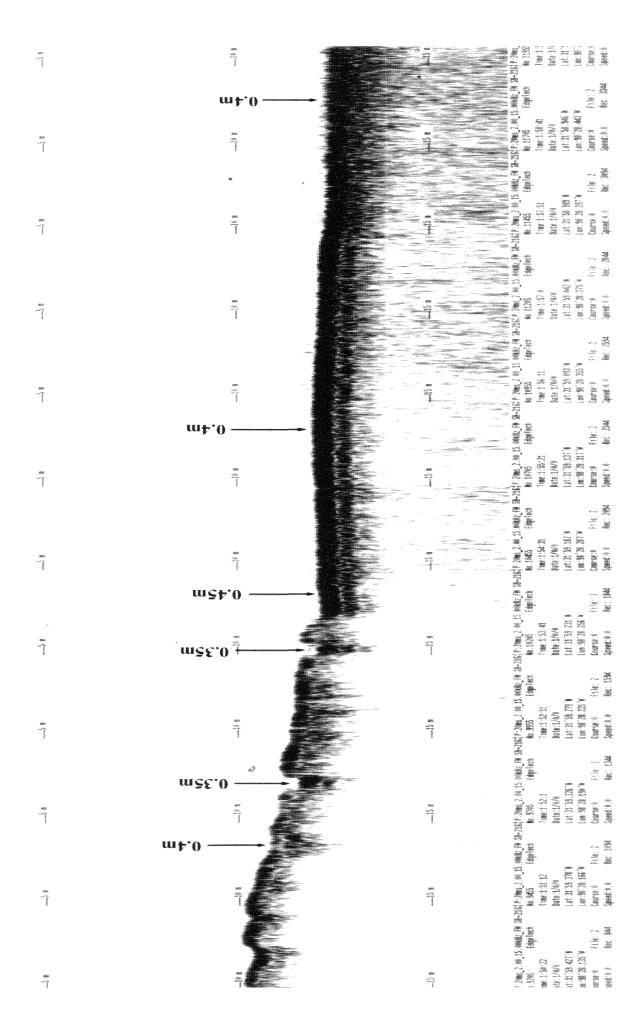


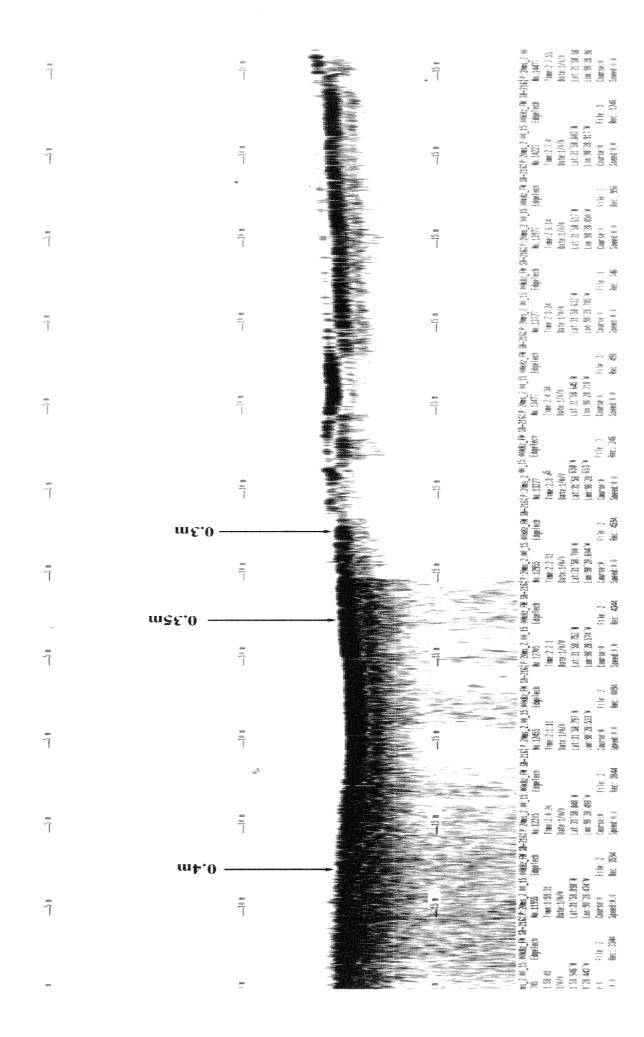


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